#### REPORT DOCUMENTATION PAGE

Form Approved OMB No. 0704-0188

Public reporting burden for this collection of information is estimated to average 1 hour per response, including the time for reviewing instructions, searching existing data sources, gathering and maintaining the data needed, and completing and reviewing the collection of information. Send comments regarding this burden estimate or any other aspect of this collection of information, including suggestions for reducing this burden, to Washington Headquarters Services, Directorate for Information Operations and Reports, 1215 Jefferson Davis Highway, Suite 1204, Arlington, VA 22202-4302, and to the Office of Management and Budget, Paperwork Reduction Project (0704-0188), Washington, DC 20505.

1. AGENCY USE ONLY (Leave blank)		3. REPORT TYPE AND DATES (	COVERED
	15 Dec 03	THE	<b>ESIS</b>
4. TITLE AND SUBTITLE		5. FUNDI	ING NUMBERS
"VALIDATION OF THE PARAMA		ONOSPHERIC	
SPECIFICATION MODEL (PRISM	<b>A)"</b>		
· · · · · · · · · · · · · · · · · · ·		1	•
6. AUTHOR(S)			•
LT ROBERT PULLIAM C	,	·	
	•		
· · · · · · · · · · · · · · · · · · ·	·	1	
7. PERFORMING ORGANIZATION NAM	/IE(S) AND ADDRESS(ES)		ORMING ORGANIZATION
OHIO STATE UNIVERSITY	•	KEPUH	RT NUMBER
		l	CTOD DOD
		İ	CI02-230
		İ	
		· .	
9. SPONSORING/MONITORING AGENC			ISORING/MONITORING
THE DEPARTMENT OF THE AIR	FORCE	AGEN	NCY REPORT NUMBER
AFIT/CIA, BLDG 125		<b>l</b> '.	
2950 P STREET		l	
WPAFB OH 45433		l	
11. SUPPLEMENTARY NOTES			
·			
		•	
12a. DISTRIBUTION AVAILABILITY STA		1401 0107	
Unlimited distribution	I EMEN I	120. DIST	RIBUTION CODE
In Accordance with AFI 35-205/AFI	Tr. C 1		
III Accordance with Art 33-203/At 11	- Sup 1	N STATEMENT A	
	Approvedici	on Unlimited	
13. ABSTRACT (Maximum 200 words)	DISTIDUM	<u> </u>	
13. Abbittadi prazimani 200 militadi		·	
	•		

20040105 018

14. SUBJECT TERMS			15. NUMBER OF PAGES
			115 16. PRICE CODE
	·	_	
17. SECURITY CLASSIFICATION OF REPORT	18. SECURITY CLASSIFICATION OF THIS PAGE	19. SECURITY CLASSIFICATION OF ABSTRACT	20. LIMITATION OF ABSTRACT

# VALIDATION OF THE PARAMATERIZED REAL-TIME IONOSPHERIC SPECIFICATION MODEL (PRISM)

# A Thesis

Presented in Partial fulfillment of the Requirements for

The Degree Masters of Science in the

Graduate School of The Ohio State University

By

Robert Charles Pulliam, B.S.

\*\*\*\*

The Ohio State University 2003

Masters Examination Committee:	Approved by	
Dr. Jeff Rogers, Advisor	Approved by	
Dr. Jay Hobgood		
, ,	Advisor  Department of Atmospheric Sciences	

DISTRIBUTION STATEMENT A

Approved for Public Release Distribution Unlimited

### ABSTRACT

The earth's ionosphere between 60km and 1000km altitude contains a significant amount of partially ionized plasma that affects the propagation of radio waves. This plasma is created when extreme ultraviolet (EUV) light from the sun strips electrons from the neutral molecules in the Earth's atmosphere. The ionosphere's free electron density is highly variable and often unstable and can adversely affect Department of Defence systems which rely on radio wave propagation. These effects include: inaccurate position readings from GPS satellites, communication disturbances, and communication outages. The Air Force Research Laboratory has developed a Parameterized Real-time Ionospheric Specification Model (PRISM) that specifies the density of free electrons in the ionosphere on a global scale. This research will focus on validating PRISM using data from GPS satellites, the Digital Portable Sounding (DPS) network, and TOPEX/Poseidon data. In order to do a complete performance analysis, several time periods (3.6 weeks) of varying solar activity will be selected. Once these periods are selected PRISM will be initialized two different ways. The first initialization will be made without any real time input data and the output will be purely PRISM climatology. As for the second initialization PRISM will be given the real time data that the Air Force Weather Agency uses and the output will be an

adjusted climatology. Once these two sets of PRISM runs are complete they will be compared to the validation data and an analysis of the improvement gained by using the input data can be made. Additionally, the performance of PRISM at different solar activities, times of the day, and varying latitudes will be explored.

Dedicated to my wife

### **ACKNOWLEDGMENTS**

I would like to express my sincere thanks to Dr. Dwight Decker (Air Force Research Lab), & Dr. Patricia Doherty (Boston College) for the data and guidance throughout this research project. They provided, advice, resources and insight that helped me tremendously throughout this research. Without their assistance this project would have not been possible. I would also like to thank my advisor Dr. Jeff Rogers for his patience and genuine interest in this project. His flexibility in allowing me to pursue a research topic outside his area of expertise is most appreciated.

I would also like to extend thanks to all the faculty in the Atmospheric Sciences Department, Dr. Rogers, Dr Hobgood, and Dr. Arnfield for their willingness to work with the time the Air Force has allowed for me to get my degree. This has truly been one of the most challenging and rewarding experiences of my life. The knowledge gained at Ohio State will help me go far in my Air Force career in being a successful weather officer.

Most importantly, I would like to thank my wife, who has gone out of her way in supporting me though this time consuming effort. Without her support this would have not been possible.

#### VITA

	Born – Duluth, Minnesota
May 1999	B.S. Physics, University of Wisconsin Commissioned in the US Air Force
June 1999	Married
1999-2002	Space Forecasting Test Manager USAF, Space Vehicles Directorate Hanscom AFB, Lexington, MA
2002-Present	Graduate Student/Air Force Officer The Ohio State University Columbus, OH

#### **PUBLICATIONS**

- 1. R. C. Pulliam, R. Rybski, "Asteroid Detection System", National Conference for Undergraduate Research, Salisbury, Maryland, Apr 1998.
- 2. R. C. Pulliam, R. Rybski, "Asteroid Detection System", National Conference for Undergraduate Research, Salisbury, Maryland, Apr 1999.
- 3. R. C. Pulliam, W. Borer, "PRISM Validation Study", Proceedings of the 2000 Space Weather Conference, Boulder CO, May 2000.
- 4. R. C. Pulliam, W. Borer, D. Decker, P. Doherty, "Operational Ionosphere Model Validation", Proceedings of the American Institute of Aeronautics and Astronautics Space 2000 Conference & Exposition, Paper #A00-42948, Long Beach, CA, Sep 2000.
- 5. R. C. Pulliam, W. Borer, "PRISM Validation Study Update", Proceedings of the 2001 Space Weather Conference, Boulder CO, May 2001.

#### FIELDS OF STUDY

Major Field: Space & Tropospheric Weather Minor Fields: Physics, Astronomy, & Math

# TABLE OF CONTENTS

Abstract	<u>Page</u> ii
110561406	
Dedication	iv
Acknowledgments	v
Vita	vi
List of Tables	xi
List of Figures	x
Chapters:	
1. Introduction	
1.1 The Ionosphere	
1.1.1 D Region	
1.1.2 E Region	
1.1.3 F Region	
1.1.4 Latitude & Local Time Variat	
1.2 Introduction to PRISM	
1.2.1 PIM	
1.2.2 The Low Latitude F Layer Mo	odel 11
1.2.3 The Mid Latitude F Layer Mo	odel 11
1.2.4 The Low & Mid Latitude Lay	
1.2.5 The High Latitude Model 1.2.6 PRISM's RTA Algorithm	
2. The Validation of PRISM	15
2.1 Past Validation	15
2.1.1 Validation Method	18
	19
	22
2.2 This Validation	
	25
	30
	32
	tudy32
2.2.5 Method	35

3. The Results	39
3.1 Local Time Analysis	39
3.2 Distance From Station	41
3.3 Kp Distributions	45
3.4 Second Kp Analysis	49
4. Summary & Recommendations	53
4.1 Summary	53
4.2 Conclusion	54
Appendix A: Sample TEC GPS Driver Data Plots for Pert	61
Appendix B: Main Analysis Program	64
Appendix C: Sample Validation File	73
Appendix D: Ten Day Distance Error Distributions	74
Appendix E: Local Time Error Distributions by Kp	81
Appendix F: Sample PRISM Output	84
Appendix G: A sample of a PRISM TEC RTA output	88
References	89

# LIST OF TABLES

<u>Table</u>		Page
1.1	Geophysical Parameter Values	9
1.2	Horizontal Grid Resolution used for Climatology	10
2.1	PRISM Altitude Profile in Kilometers	25
2.2	Input Data Quality by Day Number	28
2.3	TOPEX_Comparison6.pro Input File Sizes	36
2.4	TOPEX_Comparison6.pro Output File Variables	38
3.1	Magnetic Latitude Range for Comparisons	39
3.2	Kp Sample Distributions	52

# LIST OF FIGURES

<u>Figur</u>	<u>ce</u>	Page
1.1	Plasma Density with Height	4
1.2	Molecule Species with Height	4
1.3	Operation of PRISM	14
2.1	Conversion from Slant TEC to Vertical Equivalent TEC	16
2.2	Distribution of the Input GPS Stations for PRISM	17
2.3	Distribution of the GPS Ground Truth Stations for PRISM	17
2.4	Processed Data from the GPS Receivers	18
2.5	Station Summary Plots (PRISM vs Ground Truth).	20
2.6	PRISM Output Compared to the Input Data	20
2.7	Input data and ground truth comparisons	21
2.8	Statistics for the Ground Truth Summary Plots for PRISM	24
2.9	Distribution of GPS Stations	26
2.10	Station Quality Map	29
2.11	Sample Single TOPEX pass for Day 35 and the corresponding TEC plot	33
2.12	Data Smoothing	33
2.13	Example of TOPEX Ground Tracks for One Day	35
2.14	TOPEX_Comparison6.pro Flow Chart	37
3.1	Model Error VS Local Time	40
3.2	Distance From Station Verses Error	42
3.3	Distribution of TEC error larger than 20TEC units from	

# Chapter 1

#### Introduction

# 1.1 The Ionosphere

The ionosphere is the region of charged particles surrounding Earth between the altitudes of 60km – 1000km and is created by the ultraviolet (UV) radiation from the Sun [Jursa, 1985]. This process is called photo ionization which produces the following reactions:

$$O_2 + photon \rightarrow O_2^+ + e^-$$
  
 $O + photon \rightarrow O^+ + e^-$   
 $N_2 + photon \rightarrow N_2^+ + e^-$ 

Since the ionosphere is created by the sun's UV radiation the density is dependent on the time of day. So after sunset photo ionization stops and the free electrons begin to recombine with the  $O_2$ , O,  $N_2$ , ions [Tascione, 1994]. This recombination often eliminates some layers of the ionosphere completely. The distribution of the ionosphere is best described by the continuity equation:

$$\frac{\Delta N_e}{\Delta t} = P - L + T$$

 $-\frac{\Delta N_e}{\Delta t}$  = Rate of change in the density of electrons with time

-P = Production rate of electrons from UV radiation.

 $\cdot$ L = Loss rate of free electrons from recombination.

-T = Transport rate at which electrons are transported into or out of the volume.

The production rate (P term) is the daytime source for creating the ionosphere and during the night recombination (L term) attributes to decreasing the density of the ionosphere. The transport term represents the

movement of electrons from one area to another caused by plasma drift. Plasma drift occurs when ions move from an area of higher density to a region of lower density. This drift can come from several factors such as gravity, neutral winds, and the earth's magnetic field. The variability of this drift is dependent on the altitude because the density of the ionosphere varies with height.

The ionosphere is also considered to be a plasma. A plasma is composed of a collection of discrete ionized particles. However not every collection of charged particles qualifies as a plasma, as certain criteria must be met. Over large length scales, the medium must be electrically neutral. For a plasma composed of electrons and protons, electrons are attracted toward protons and repulsed from other electrons by electrostatic forces. So at a certain distance from a charged particle, its charge can no longer be seen due to the shielding, or screening, of the other charged particles around it. This distance is called the Debye length (l<sub>D</sub>), which is defined by the following equation:

$$l_D^2 = kTe_o/nq^2$$

 $-l_D = Debye length$ 

-k = Boltzman constant

T = Temperature

 $-e_0$  = the permittivity of free space

-n =the plasma number density

-q = the unit electric charge.

Within a sphere of this radius, there are

$$N_D = 4 \pi n l_D^3$$

other charged particles. An ionized gas is termed a plasma when:

$$g = 1/N_D < < 1$$

·g is called the plasma parameter.

This plasma parameter depends on the charged particle density and the average energy of the particles measured by the temperature. In the ionosphere, g ranges from  $10^{-4}$  to  $10^{-6}$ . So considering the ionosphere a plasma is certainly valid.

Earth's ionosphere is divided into several regions designated by the letters D, E, F (Figure 1.1). These regions may be further divided into several regularly occurring layers, such as F1 and F2. Historically, these divisions arose from the successive plateaus observed in the electron number density. Distinct ionospheric regions develop because (1) the solar spectrum deposits energy at different heights depending on absorption characteristics of the atmosphere, (2) the physics of recombination depends on the density, which exponentially decreases with height, and (3) the composition of the atmosphere changes with height. Thus the main ionospheric regions can be associated with different governing physical processes. In figure 1.2 you can see the dominating molecules as a function of height [Canck, 2002].

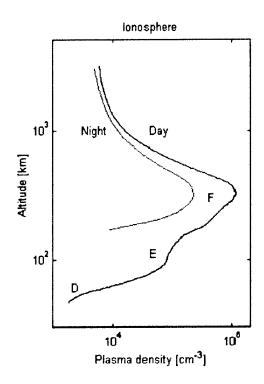


Figure 1.1: Plasma Density [cm<sup>-3</sup>] with Height [km].

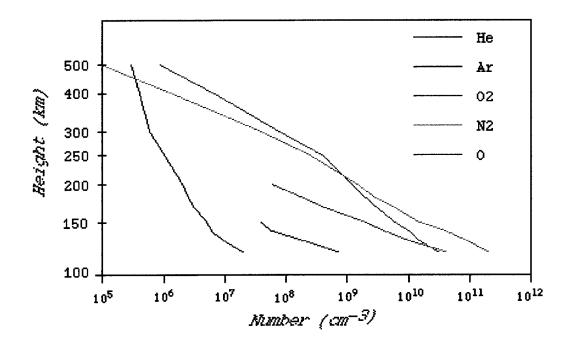


Figure 1.2: Molecule Species with Height

# 1.1.1 D Region

The D region is the region of the ionosphere that is approximately 60km to 90km above the surface of the earth. The charged particle number density ranges from between 10<sup>2</sup> to 10<sup>4</sup> cm<sup>-3</sup> during the day and completely vanishes during the night due to recombination. During sunrise ions are formed by the ionization of atmospheric neutrals by the Extreme Ultra Violet (EUV) radiation (.1-.11um). Due to the relatively high ambient atmospheric pressure, many negative ions are produced by electron attachment to atomic and molecular neutrals. Positive and negative ions of N<sub>2</sub>, O<sub>2</sub>, and O are dominant constituents.

#### 1.1.2 E Region

The E region lies between an altitude of 90 and 150 km above the earth. The electron densities range on average between 10<sup>5</sup> cm<sup>-3</sup> (in the daytime) to 10<sup>4</sup> cm<sup>-3</sup> (at night). Ions in this region are mainly O, O<sub>2</sub>+, and N<sub>2</sub>+, formed by EUV radiation below .11um. Other subdivisions, isolating separate layers of irregular occurrence within this region, are also labeled with an E prefix, such as the thick layer, E2. Unlike the D region, the E region can persist throughout the night as a result of dense patches of ionization called Sporadic E. Upward propagating gravity waves and tides collect this ionization into thin, downward propagating layers of enhanced ionization which can occur at all latitudes [Hargreaves, 1992]. The presence of the nighttime E region and sporadic E are thought to be due to electron and meteor bombardment from space.

# 1.1.3 F Region

The region above about 150 km is known as the F region. This region is often divided into the F1 and F2 regions. The F1 region (150km to 200km) has a maximum electron number density of a few times  $10^5$  cm<sup>-3</sup> at about 200 km altitude, and the density of the F2 region varies between  $10^6$  and  $10^5$  cm<sup>-3</sup> between day and night respectively. The altitude of maximum electron density in the F2 region is roughly 350 km. This peak is highly variable depending upon daily, seasonal, and sunspot-cycle variations. It is important to note that even at this maximum, the charged particle number density is less than the number density of neutral atmospheric gas; electron density ~  $10^6$  cm<sup>-3</sup>, while neutral number density ~  $10^9$  cm<sup>-3</sup>. The F region is formed by ionization of atomic oxygen by Lyman emissions and by emission lines of He. In the lower part of the F region, O+ ions readily transfer charge to neutrals forming N<sub>2</sub>+. In the F2 region, O+ remains the dominant ion.

# 1.1.4 Latitude and Local Time Variations

Not only does the ionosphere vary with altitude, but it also varies with latitude, local time, and season. These variations are largely dependent on solar zenith angle, which controls the amount of photoionization (appendix G contains a sample PRISM TEC output for 24 hours). The response of the ionosphere to sunrise and sunset is very rapid; electron densities can increase/decrease by two orders of magnitude within an hour of dawn/dusk, revealing the dependencies on zenith angle. A dependence on latitude is because as the northern hemisphere enters into its winter months and the solar zenith angle increases, the photoionization rate drops resulting in a

decrease in electron density. Additionally, ionization due to particle precipitation demonstrates electron content dependence on latitude. Particle precipitation happens when charged particles follow the earth's magnetic field lines and bombard the high latitude ionosphere resulting in large electron production rates in the high latitudes.

Other phenomena that influence the electron content of the ionosphere are the Sudden Ionospheric Disturbance, the Appleton Anomaly, and the Midlatitude F2 Winter Anomaly. The Sudden Ionospheric Disturbance occurs within minutes of a strong solar flare. The X-rays emitted from the flare lead to a large ionization of the ionosphere and is normally about an hour of duration. The Appleton Anomaly (or Equatorial Anomaly) is high concentrations of electrons on either side of the geomagnetic equator in the post sunrise sector (+/- 15 degrees magnetic latitude). The Mid-latitude F2 Winter Anomaly is the transport of plasma by neutral winds from the summer hemisphere to the winter hemisphere causing the daytime F2 peak electron density in wintertime to be as much as four times greater than the density in the summertime hemisphere. The Winter Anomaly occurs during solar maximum conditions and is most apparent between 45-55 north magnetic latitude.

The eleven-year solar cycle also influences the ionosphere. During periods of high solar activity there is an increase in the number of high energetic particles discharged from the sun, which increases particle precipitation. Particles precipitating can increase ionization by an order of magnitude from solar minimum to solar maximum. Also during periods of

high solar activity there is an increase in the number and strength of geomagnetic storms. However geomagnetic storms are observed during quiet solar conditions too [Tascione, 1994].

#### 1.2 Introduction to PRISM

PRISM is comprised of two components, a Parameterized Ionospheric Model (PIM), and a Real Time Adjustment (RTA) algorithm [Daniell and Brown, 1995]. PIM is a global specification of ionospheric free electron densities that was created from curve fits to the output fields of several theoretical ionospheric models. This description of the ionosphere's average behavior is referred to as climatology. The inputs for PIM are: solar activity (F10.7) which is a measure of the sun's radiation at a wavelength of 10.7cm, magnetic activity (Kp) which is an indicator of the general level of magnetic field strength variations measured at the Earth's surface, and the strength of the sun's interplanetary magnetic field (IMF) vectors B<sub>v</sub> and B<sub>z</sub> measured by a Magnetometer aboard the Advanced Composition Explorer (ACE) satellite. The output of PIM is a global 3-dimensional specification of electron densities which includes the ionosphere's E-region (~90-155km) and the F2-region (~250-1000km), critical frequencies of the E and F2 region (foF2) where the critical frequencies are the highest HF frequency able to be reflected by the layer without penetrating it, and heights of the E layer and F2 layer. Variations in the critical frequency are caused by variations in the density of the ionosphere. A larger density (or higher TEC value) will reflect a higher wavelength than smaller density. As for the output, the electron densities from PIM and PRISM are available in two formats: (1) gridded output on a

regional or global grid in geographic or geomagnetic latitude and longitude,
(2) output at a set of user specified points in geographic coordinates on the
Earth's surface.

#### 1.2.1 PIM

The parameterised database was generated using an ensemble of four separate physical models: (1) a low latitude F layer model (LOWLAT), (2) a mid latitude F layer model (MIDLAT), (3) a combined low and middle latitude E layer model (ECSD), and (4) a high latitude E and F layer model (TDIM). All four models are based on a tilted dipole representation of the geomagnetic field and a corresponding magnetic coordinate system. In addition, information on heat transport, thermospheric winds and plasma drift velocities were incorporated into these models. From running these four models a parameterised representation of the ionosphere database was developed. This process took into account different geomagnetic conditions (Kp), solar activity (F10.7), universal times, Interplanetary Magnetic Field (IMF) By direction, geomagnetic latitudes and longitudes, and days of the year. Since it would take a long time and a large amount of computer memory in order to account for every possible combination of these parameters these models were run for a relatively small number of possible conditions [Daniell and Brown, 1995]. This parameterisation was accomplished in a two-step process. First, the models were used to generate a number of databases for a discrete set of varying solar activity. Each of these databases consisted of ion density profiles on a discrete grid of latitudes and longitudes for a twenty-four hour period. Second, in order to reduce the

storage requirements the databases were approximated with semi-analytic functions. The tables below shows the values used in generating these databases as well at the grid resolution used.

Model	Solar activity F10.7	Magnetic activity Kp	IMF By	Day of the year	Databases
LOWLAT	70, 130, 210	N/A	N/A	80, 172, 264, 355	36a
MIDLAT	70, 130, 210	1, 3.5, 6	N/A	80, 172, 264, 355	54b
ECSD	70, 130, 210	1, 3.5, 6	N/A	80, 172, 264, 355	54c
TDIM	70, 130, 210	1, 3.5, 6	+,-	80, 172, 264, 355	324d

- a. 3 seasons X 3 solar activities X 4 longitude sectors
- b. 3 seasons X 3 solar activities X 3 magnetic activities X 2 hemispheres
- c. 3 seasons X 3 solar activities X 3 magnetic activities X 2 species
- d. 3 seasons X 3 solar activities X 3 magnetic activities X 2 By's X 3 species X hemispheres.

Table 1.1: Geophysical Parameter Values

Model	Magnetic Latitude	Magnetic Longitude	UT (Magnetic
			Local time [MLT])
LOWLAT	-32 to 32 in 2 deg	30, 149, 250 and	MLT: 0-23.5 in .5
	steps	329	hour steps
MIDLAT	30 to 74 and -30 to	0 to 345 in 15 deg	MLT: 1-23.5 in 2
	-74 in 4 deg steps	steps	hour steps
ECSD	-76 to 76 in 4 deg	0 to 345 in 15 deg	MLT: 1-23.5 in 2
	steps	steps	hour steps
TDIM	51 to 89 and -51 to	Magnetic Local	MLT: 1-23.5 in 2
	89 in 2 deg steps	Time .5- 23,5 in 1	hour steps
		hr steps	

Table 1.2: Horizontal Grid Resolution used for Climatology

# 1.2.2 The Low Latitude F Layer Model

The low latitude F region model (LOWLAT) was originally developed by Anderson, [1973]. It is designed to solve the diffusion equation for O+ along a magnetic flux tube. Normally, the entire flux tube is calculated with chemical equilibrium boundary conditions at both feet of the tube. A large number of flux tubes must be calculated in order to build up an altitude profile. Since heat transport is not included in this model, ion and electron temperature models must be used. The LOWLAT model makes use of the of the ion and electron temperatures by a model developed by Brace and Theis [1981] and the Horizontal Wind Model (HWM) of Hedin [1988] for the thermospheric winds.

The critical feature incorporated in the low latitude model is the dynamo electric field. The horizontal component of this field drives upward convection due to the earth's electric and magnetic field, and this can significantly modify profile shapes and densities. This phenomenon is responsible for the equatorial anomaly, crest in the ionisation on either side of the magnetic equator at ±15 to 20 degrees magnetic latitude. In the current version of PRISM the ExB vertical drift used for these calculations was based on the empirical models derived form data from the Atmospheric Explorer-E satellites [Fejer, 1995]. Which are consistent with the drifts measured at Jicamarca, Peru but include longitudinal variations as well.

### 1.2.3 The Mid latitude F Layer Model

The mid latitude F region model (MIDLAT) is the same as the low latitude version, except that the dynamo electric field is not included.

Complete flux tubes are followed, but neither horizontal nor vertical convection is included. The computer resource requirement of MIDLAT are far less that those of LOWLAT. As long as the boundary between low and middle latitudes is chosen so that the electric field is negligible on the boundary flux tubes, the two models give identical results at the boundary ensuring continuity across that boundary. For the PRISM development the same temperature model [Brace and Theis, 1981] and the same thermospheric wind model [Hedin, 1988] were used.

## 1.2.4 The Low and Midlatitude E Layer Model

The low and mid-latitude E region model (ECSD) was developed by

Dwight T. Decker and John R. Jasperse and incorporate photoelectrons

calculated using the continuous slowing down (CSD) approximation

[Jasperse, 1982]. Ion concentrations are calculated assuming local chemical equilibrium. A small nighttime source is included to ensure that an E layer is maintained throughout the night.

#### 1.2.5 The High Latitude Model

The high latitude model (incorporating both E and F layers) is the Utah State University (USU) Time Dependent Ionosphere Model (TDIM). This model is similar to the low and middle latitude models except that the flux tubes are truncated and a flux boundary condition is applied at the top. In addition, the flux tubes move under the influence of the high latitude convection electric field. In the low latitudes, because the magnetic field is mainly horizontal, the effect of the electric field is mainly vertical, and the electric-field-driven convection is horizontal. TDIM includes an E layer model that

incorporates the effects of ionisation by precipitating auroral particles. The ion production rates used were calculated using the B3C electron transport code [Strickland, 1994] and incident electron spectra representative of DMSP SSJ/5 data. The characteristics of the electron spectra were taken from the Hardy [1987] electron precipitation model.

# 1.2.6 PRISM's RTA Algorithm

The availability of real time data permits operation of PRISM's RTA algorithm that adjusts the model output to fit the real time measurements. The real time data include: density profile parameters (foF2=density of F2 layer, hmF2=height of the F2 layer, foE=density of E layer and hmE=height of the E layer), total electron content (TEC) which is the line integral of electron density from the receiver to the satellite (1 TEC unit =  $10^{16}/m^2$ ), and a variety of in situ plasma and precipitating particle measurements from the Defence Meteorological Satellite Program (DMSP). The amount the RTA algorithm modifies the PIM output profiles to match the data is determined by the following weighting function:

 $W(lat,lon)=1/d_n(lat,lon) \qquad \qquad d_n(lat,lon)=.5[1\text{-}cos\Upsilon_n(lat,lon)]$  Where  $d_n$  is the distance between the data point and the point to be adjusted in PIM, lat and lon are the latitude and longitude, respectively, and  $\Upsilon$  is the angle of separation between the two points with respect to the center of the Earth. Figure 1.3 below show the data flow in the operation of PRISM.

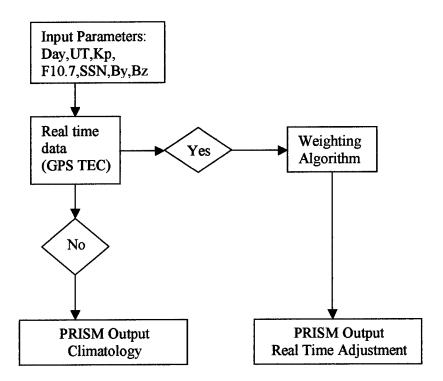


Figure 1.3: Operation of PRISM

# Chapter 2

#### The Validation of PRISM

#### 2.1 Past Validation

A previous validation of PRISM was conducted by Pulliam [2000] and focused on determining the improvement gained when running PRISM with real time GPS data. In this study data were processed and quality controlled from sixty-two dual frequency GPS stations from the International GPS Service (IGS) network for a one week period of quiet magnetic activity (13 Jan 00 - 19 Jan 00). Thirty-seven of these stations were selected as inputs for PRISM while the remaining twenty-five were used as ground truth for the validation. By using the Receiver Independent Exchange (RINEX) data from the GPS receivers the Total Electron Content (TEC) values were calculated from the differential group delay and phase advance measurements routinely made by the receivers which monitor the L1 (1575.42 MHz) and L2 (1227.6 MHz) frequencies of GPS satellites. Since the GPS signal is a time-encoded transmission, the time of flight for each of these signals can be calculated. By comparing these two time of flight measurements it is then possible to calculate the refraction of the GPS signal caused by the ionosphere. This result is then converted to a slant TEC measurement along the line of sight (LOS) from the receiver to the satellite. Where a TEC unit is the line integral of electron density from the receiver to the satellite (1 TEC unit =  $10^{16}/\text{m}^2$ ). However, in order to use these measurements for PRISM inputs, they must first be converted to vertical equivalent TEC (VETEC). This conversion assumes that the height of the ionosphere is 400km. The formula for this

conversion is: VETEC = (Slant TEC) x cos[arcsin(.94092 x cos(Elevation angle))]. Where the elevation angle is the angle from the horizon to the satellite. Figure 2.1 shows the geometry of this conversion.

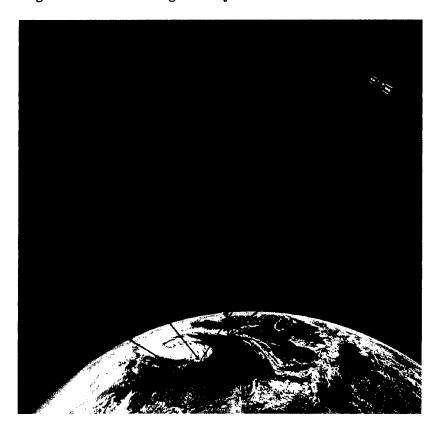


Figure 2.1:Convertion from Slant TEC to Vertical Equivalent TEC (VETEC)

The distribution of the IGS stations used in this study are represented in Figures 2.2 and 2.3. The triangles represent the location of the input stations and the asterisks represent the location of the ground truth stations.

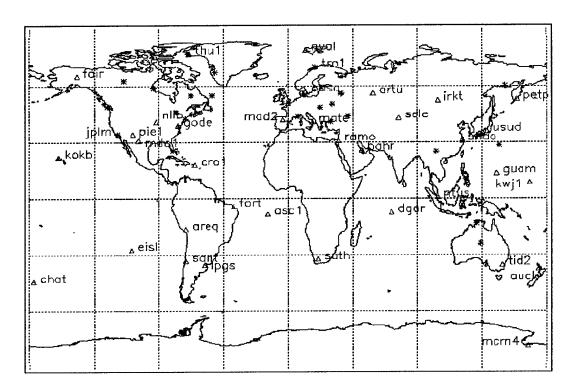


Figure 2.2: Distribution of the Input GPS Stations for PRISM.

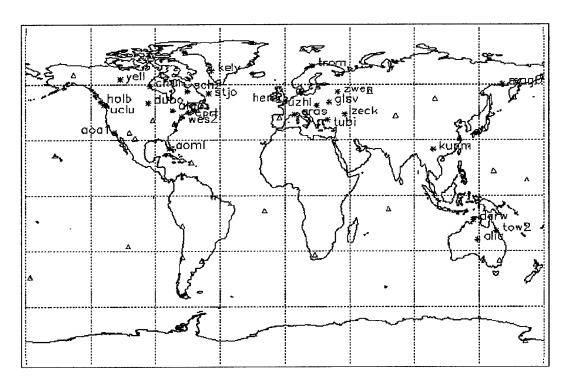


Figure 2.3: Distribution of the GPS Ground Truth Stations for PRISM.

#### 2.1.1 Validation Method:

PRISM was run twice for the 1 week period at hourly intervals to produce two different global 3-dimensional ionospheric specifications. The first run was done with no input data and produced a PIM output based only on climatology. In the second run, PRISM was given the TEC data from the thirty-seven input stations to produce an adjusted climatology output. Figure 2.4 is an example of the TEC data used and shows its diurnal pattern. The variation in the individual station maximum is due to the stations magnetic latitude. For example station AREQ which is in Peru is peaking near 100 TEC units, and ARTU at a higher latitude is around 40 TEC units.

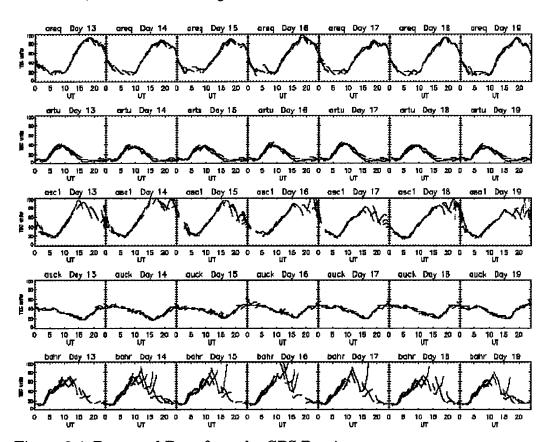


Figure 2.4: Processed Data from the GPS Receivers.

For the validation the ground truth slant TECs (STEC) were compared to PRISM's STECs. The PRISM STEC was calculated by integrating along the ground truth LOS through PRISM's 3-dimensional electron density specification from an altitude of 90km to 1600km. This comparison was done on both PRISM runs for all STEC values for all ground truth stations. From these results it was determined how much the input data improved the accuracy of the model with respect to the ground truth.

#### 2.1.2 The Results.

The major objective of this study is to compare the PRISM STEC (model value) with the ground truth STEC (GPS measured value). Figure 2.5 shows the difference between (PRISM's STEC) minus (Ground truth STEC) vs number of occurrences (individual measurements). The errors are expressed in TEC units. The solid line represents the PRISM runs with no input data (climatology), and the dashed line represents the PRISM runs with the input data. Each of these station's summary plots show a significant improvement when PRISM is given the thirty-seven stations of data. Both the large positive and negative errors were reduced and there is a larger number of these differences centered around zero. Figure 2.6 shows the PRISM outputs compared with the input slant TEC data. This was done to gain confidence in PRISM's ability to assimilate the real time data and to assess how well PRISM was able to reproduce the slant TEC when given VETEC.

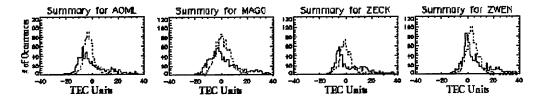


Figure 2.5: Station Summary Plots (PRISM vs Ground Truth)

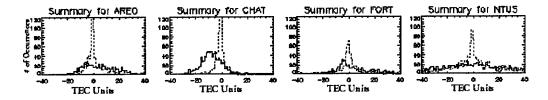


Figure 2.6: PRISM Output Compared to the Input Data.

As seen from this diagram the model does a very good job at assimilating the data. However a closer look was taken on this data because a perfect assimilation would represent a delta function. When doing this it was noticed that 90% of the larger error came from data that had a very low elevation angle, which is the angle of the GPS satellite with respect to the horizon. This error is thought to be caused by multi-path interference caused by obstacles on the surface. So to eliminate this error from this study the elevation was limited to 45 degrees and greater. In Figure 2.7 all the summary plots of the stations were combined in one concise plot. The addition of input data into PRISM results in an overall improvement in the error distribution. However, the detail as to how the individual ground truth stations are affected is lost in this display.

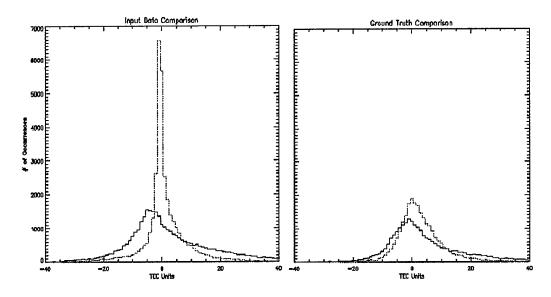


Figure 2.7: Input data and ground truth comparisons.

The dashed line represents the PRISM run with input data.

The solid line represents the PRISM run with no input data

# 2.1.3 Summary and Conclusion

In this study more than 150,000 measured STECs were made from sixty-two globally distributed IGS receivers during this one week period in January. Thirty-seven of these receivers were used as inputs in PRISM. The remaining twenty-five, were set aside to be used as ground truth. In this validation effort PRISM was run two ways:

Run 1 = PRISM run with no input data (climatology).

Run 2 = PRISM run with the thirty-seven stations (RTA).

In the first PRISM run, the only model inputs were the date, F10.7, Kp, and By and Bz components. However for the second PRISM run the GPS vertical TEC data were also given to PRISM to do the RTA. With these two sets of runs a comparison was made to determine the improvement when input data was used. The results of this comparison showed an overall improvement when the RTA was made to the input data. Figure 2.8 summarizes the TEC specification improvements in TEC units for each GPS receiver.

Overall, the average standard deviation of the TEC error when compared to the ground truth data was reduce by 6.1 TEC units (44% improvement over climatology), the mean error was reduced by 2.3 TEC units (39% improvement over climatology), and the root mean square (RMS) error was reduce by 6.4 TEC units (42% improvement over climatology) during this validation time period. However, the most striking results came from the station by station analysis (Figure 2.8). For the majority of the mid latitude station the errors were usually

cut in half with the use of the input data, although a few stations showed little to no improvement. By looking at these three stations closer (KELY, YELL, TROM) three conclusions were drawn. The first was that the model with no input data did fairly well in the station's region and input data did not really have an effect. Secondly, the model is slightly misplacing a TEC gradient, or that the slant to vertical conversion is either over or under estimating the input VETEC which would incorrectly adjust the model. Lastly is the fact that these stations are all in the high latitude region. This region of the ionosphere is highly dynamic which makes it very difficult to model. So PRISM may not be the first choice in models for high latitudes. For this region a physics based model may be more appropriate.

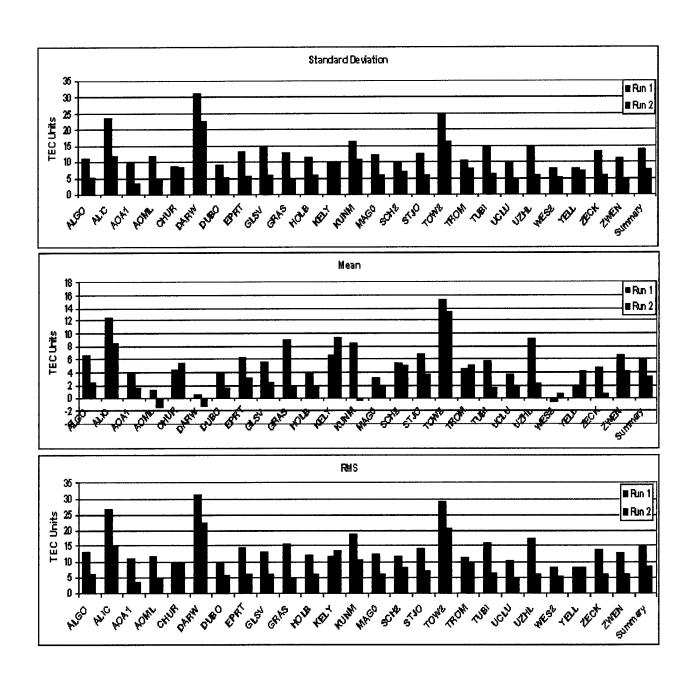


Figure 2.8: Statistics for the Ground Truth Summary Plots for PRISM.

# 2.2 This Validation

## 2.2.1 Introduction

In this validation PRISM was set up to run hourly to produce electron densities on a global grid (a partial PRISM output is in appendix F). PRISM has a maximum resolution of two degrees by two degrees. However, in order to reduce the computation time and hard drive space, the resolution was set to four degrees by four degrees. For each of the grid points PRISM computes the electron density at fifty different altitude levels. Table 2.1 below shows the altitudes used for this study.

90	160	290	750
95	170	300	800
100	180	320	850
105	190	340	900
110	200	360	1000
115	210	380	1100
120	220	400	1200
125	230	450	1300
130	240	500	1400
135	250	550	1500
140	260	600	1600
145	270	650	
150	280	700	

Table 2.1: PRISM Altitude Profile in Kilometers

A finer altitude resolution was used for the lower levels of the ionosphere due to the greater variability in this region. The lower ionosphere has the greatest electron density. PRISM was run twice at hourly intervals during the period of February 4 2002 (day 35) – July 19 2002 (day 200). The first run was done with no input data (only Kp, Day, UT, F10.7, SSN, By and, Bz) which produced a PRISM output based only on climatology (CLM). In the second run, PRISM was given the GPS TEC data used by the Air Force Weather Agency (AFWA) to run PRISM in real time to produce an adjusted climatology output (RTA). This GPS data are provided by the Jet Propulsion Lab in Pasadena CA from a sub set of their global network of GPS stations operating in real time. The figure below shows the locations of these GPS stations.

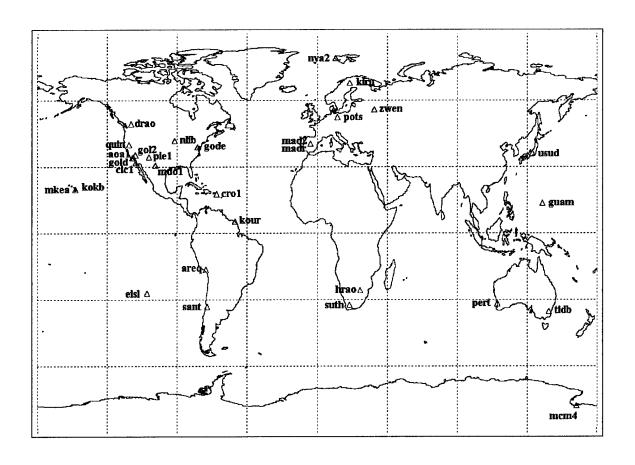


Figure 2.9: Distribution of GPS Stations

The first step in the validation is to look at the quality and consistency of the GPS data given to PRISM. To do this the TEC data were plotted for each station for every day. This resulted in 165 daily plots for each of the thirty stations. Appendex A contains a sample of these plots for station PERT which is located in Australia. A summary of the station inconsistencies is listed in table 2.2.

Findings		
Excellent, S: 55, 90,106 M: 57		
Fair, S:35-72,77 M:44-46,64,89,90		
Good, S:56,60,65,69,77,106,126,141,157,186 M:48,200		
Excellent, S: 65,77,106,153,187,186 M: 200		
Good, S:46,65,77,61,62,106,125,153,157,167,186		
Poor - many days very sparse, M:57-59,60,134-143,147,148,187,200		
Poor - many days very sparse, M: 47,56,69,81,82,90-106,113,124-129,166,200		
Excellent, S:77,106,107,186 M:		
Excellent, S:106,153,167,186		
Good - data other than gaps, S:65,77,142,143,154 M:130-141,147-153,166,185-200		
Good, S:65,77,106,127,167,186 M:200		
Excellent, S:106,186		
Fair, S:52,65,77,82-84,106,153,157,167,178,186,189 M:61,62,81,132		
Many days very sparse, M:35,37,41,62,81,86,91-93,131,132,167,168		
Excellent, S:73,74,106,157,186		
Excellent, S:65,77,106,153,157,186		
Fair - many days very sparse, M:41,179-181,200 (not to bad for high lat station)		
Excellent, S:77,106,129,162 M:163,186,189,200		
Excellent, S:58,77,106,153,157,167,180,186 M:200		
Good, S:65,77,99-102,106,111,112 M:200		
Excellent, S:77,106,107,157, M:90-92		
Excellent, S:50-52,61-63,106,107,132,157,178,186-189 M:118		
Excellent, S:45,70,77,78,106,157,186 M:69,200		
Excellent, S:65,77,81,82,106,153,157,167,186,189		
Fair many missing days, S:46, 77,78,102,134,153,154 M:48-63,103-133,166-200		
Excellent, S:38,39,55,77,106,115,153,157,158,167,186		
Poor, M:58-90		
Excellent, S:65,77,98,106,135,153,157,186		
Good,S:46,65,77,82,106,157,187,186 M:48-54		
Excellent,S:51,65,77,61,62,106,121,125,183,157,187,186		

Table 2.2: Input Data Quality by Day Number (S=Sparse data, M=Data Missing)

From this analysis it was determined that the overall quality of the data was quite good. There are no large outlying TEC data values that would cause the RTA algorithm to falsely adjust the model. However there were a few days

[77,106,157,186] where most of the stations had very sparse data. The sparseness of the data would produce a PRISM output file very similar to the climatology output file and would affect the results of the validation. So these days were removed from the validation. Also, there was a few days that most of the stations didn't have data at all and they are as follows: [42, 47, 55, 56, 68, 75, 76, 166, 185, 200]. These days were also removed. Figure 2.10 shows the data quality spatially for the remaining days used in this study. The station locations are color coded to correspond to the data quality for each individual GPS receiver.

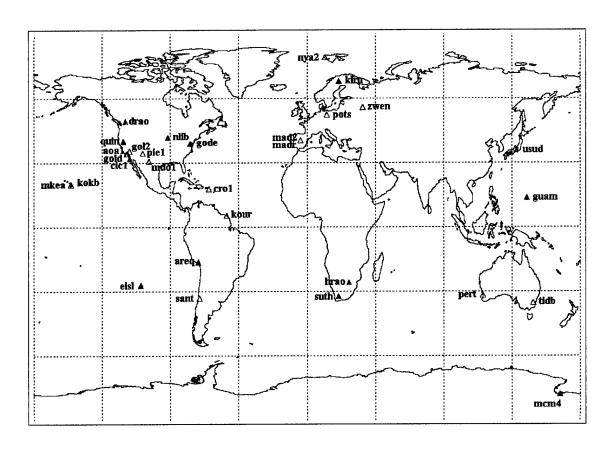


Figure 2.10: Station Quality Map (Blue=Excellent, Green=Good, Fair=Purple, Red=Poor)

With these two sets of PRISM runs (CLM &, RTA) a comparison will be made to the TOPEX validation data. By comparing the RTA comparison to the CLM comparison an assessment can be made of the improvement gained by using the GPS input data. This study will also greatly expand upon the past validation, which only looked at the improvement of the overall error when running the model with GPS data for a week of quiet solar activity. It should also be noted that the previous validation data were limited primarily to the mid latitude region. However, this study does a more detailed study by examining the errors as a function of magnetic latitude (low, mid, and some high latitudes) and local time and how they vary during periods of differing solar activities. By doing this a performance analysis can be made of PRISM which will be invaluable to both the forecaster and the model developer. By knowing the strengths and weaknesses of PRISM the forecaster can apply the appropriate confidence level to the model output. This research will also provide the Air Force Research Lab model developers with the insight to make improvements to the next version of PRISM and establish a bench mark for the future generation of ionospheric models.

### 2.2.2 Validation Data

The data used for the validation are from the TOPEX/Poseidon mission (1992) which is a joint endeavor between NASA and the French space agency, Centre National d'Etudes Spaatiales (CNES), designed to study global ocean

dynamics. This satellite has an orbit of 1336 km, and an inclination of 66 degrees. The time to complete one orbit is approximately 112 minutes, and it has an orbital speed of 7.2 km/s. This provides world wide (over-ocean) coverage of vertical TEC within a longitude range of 0 to 360 degrees and a latitude range of –66 to 66 degrees. There are 127 revolutions in a TOPEX cycle, which covers the same surface tracks every 10 days. The instrument used to get the TEC data is a dual frequency altimeter (C Band 5.3 GHz, Ku Band 13.6 GHz) that takes measurements at a rate of one per second. The measurement range is given by:

$$R_{measured} = R_{true} + \Delta R_{ionosphere} + \Delta R_{other}$$

where  $R_{true}$  is the true range,  $\Delta R_{ionosphere}$  is the ionospheric range error at the frequency C or Ku, and  $\Delta R_{other}$  are the range errors due to other frequency dependent and non frequency dependent sources. The range error,  $\Delta R_{ionosphere}$  (centimeters), has the form  $\frac{b_C}{f_C^2}$  or  $\frac{b_{Ku}}{f_{Ku}^2}$  where  $b_i$  equals 40.3  $TEC_{vertical}$  with  $f_i$  expressed in gigahertz. The measured range equations for C and Ku bands provide an expression for the differential ionospheric correction:

$$\Delta R_{ionosphere} = [40.3 TEC_{vertical}][(f_{Ku}^2 - f_C^2)/f_C^2 f_{Ku}^2]$$

The vertical TEC follows as:

$$TEC_{vertical} = \Delta R_{ionosphere} [f_C^2 f_{Ku}^2 / (f_{Ku}^2 - f_C^2) 40.3]$$

One TEC unit (TECU) ( $10^{16} \ el/m^2$ ) corresponds to 12.17mm range error at the given TOPEX altimeter frequencies [Vladimer, 1999].

#### 2.2.3 TOPEX Limitations

The TOPEX data have some limitations when being used to study vertical TEC and they are as follows.

- A study done by Callahan [1993] and Imel [1994] estimated the error in this data to be about 3 TECU. However in this study the TOPEX data will be regarded as truth.
- Another limitation is that the data do not obtain measurements over land.
   This will produce large gaps in coverage over the continents.
- 3. Because of the 66 degree inclination of the orbit a large portion of the high latitude is not sampled. However for this study the major focus is the low and mid latitudes so this shouldn't be an issue in this research.

# 2.2.4 TOPEX Data Used for This Study

TOPEX data from February 4<sup>th</sup> 2002 (day 35) – July 19 2002 (day 200) were provided by Dr. Patricia Doherty at Boston College. The raw altimeter data are first converted to vertical TEC data as outlined in the previous discussion. The

figure below shows an example of a single pass that is for February  $4^{th}$  (Day 35) from UT 20:56 to UT 21:49 along with a PRISM plot of the TEC at 2100 UT.

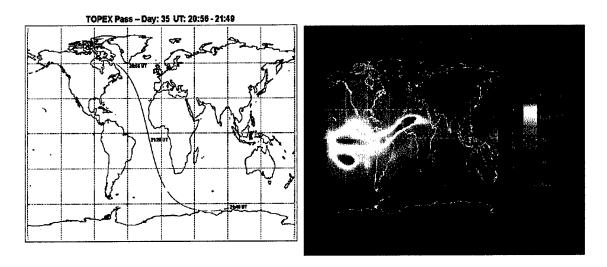


Figure 2.11: Sample Single TOPEX pass for Day 35 and the corresponding TEC plot.

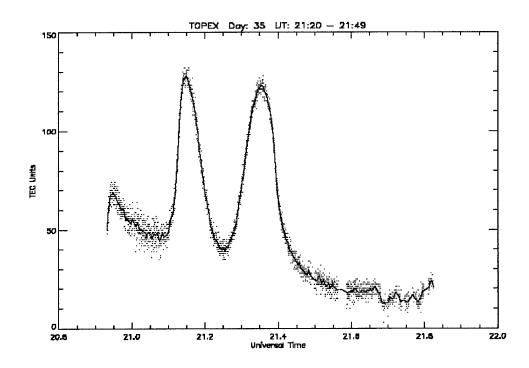


Figure 2.12: Data Smoothing (Black=1 second data, Blue=12 second average)

Figure 2.12 is the plot of the TEC for this TOPEX pass. This figure demonstrates the smoothing that was done in order to reduce some of the noise in the TOPEX data. For the smoothing, the 1 second data are averaged over 12 seconds, which converts the point measurement of vertical TEC to a 12 second average TEC value. This averaging corresponds to a distance of 86.4 km. The latitudes and longitudes of this 1-second data were also averaged to produce a corresponding 12-second average position. Since the PRISM resolution is only 4x4 degrees this averaging will not have an effect in the results and will greatly aide in reducing the number of comparisons. Finally, to get an idea of the TOPEX data coverage the ground track for day 35 from 0000 UT to 2359 UT was plotted on a world map (Figure 2.13).

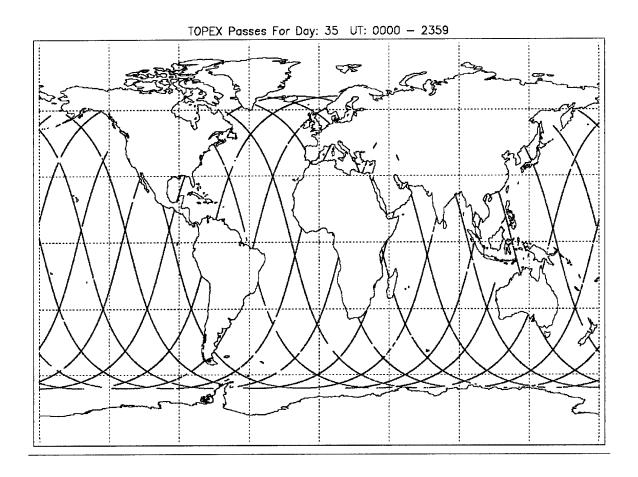


Figure 2.13: Example of TOPEX Ground Tracks for One Day

# **2.2.5 Method**

For the actual validation the two sets of PRISM runs will be organized into two 3 dimensional arrays to assist in the comparison to the TOPEX data. One array will be for the climatology run (CLM) and the other will be for the real time adjustment run (RTA) with GPS data. The dimensions of each array will be longitude (180), latitude(90), UT(24), and the corresponding value will be TEC. Then an interpolation will be made along these three dimensions for every 12-second point of TOPEX TEC data and the PRISM TEC value will be obtained.

Also, by using the latitude and longitude of the TOPEX data the distance to the closest driver station will also be calculated. This comparison looks at the GPS input file used in the RTA run to determine which stations were present at that time to do the distance calculation. In order to automate this comparison an IDL program was written called TOPEX\_Comparison6.pro (See Appendix B). This program takes one day of data and creates an output file containing the comparison data. Figure 2.14 shows how this program does the comparison. To get an idea of the amount of data processed by this program, Table 2.3 lists the data used for one day of data (day 35). The output of this program is one file containing all the TOPEX comparisons for one day, and it is this file that will be used for the error analysis. These daily files are on average only .45 [Mb]. So over the entire validation period this program takes 12045 files (20349MB) and converts it to 165 (74.25MB) files for the analysis. In appendix C you can see a sample page of one of these files. Table 2.4 shows the content contained in each of these 165 daily comparison files.

Input Files	# of Files	Avg. Size [Mb]	Total Size [Mb]
PRISM CLM Files	24	2.557	61.368
PRISM RTA Files	24	2.557	61.368
PRISM Driver Data	24	.015	.36
TOPEX Data	1	.230	.23
TOTAL	73	5.359	123.326

Table 2.3: TOPEX\_Comparison6.pro Input File Sizes.

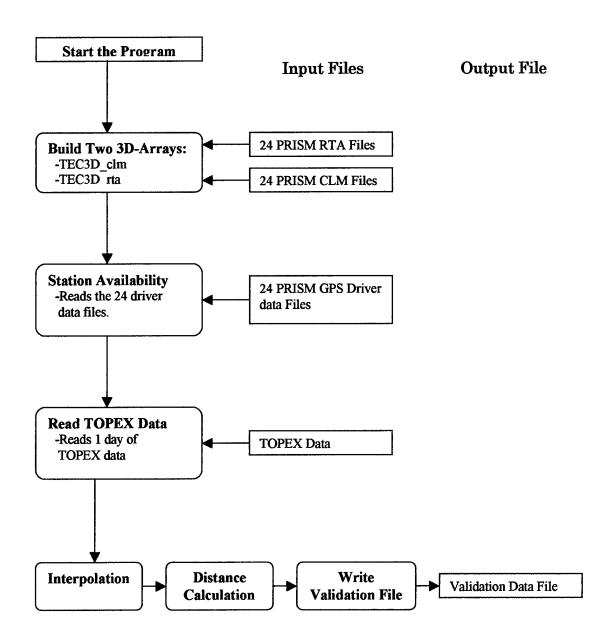


Figure 2.14: TOPEX\_Comparison6.pro Flow Chart

Variable	Discription		
Day	-Day of observation.		
UT	-Universal time at the point measurement.		
LT	-Local time at the point measurement.		
Lat	-Latitude of the point measurement.		
Lon	-Longitude of the point measurement.		
Mag lat	-Magnetic latitude of the point measurement.		
Mag Lon	-Magnetic longitude of the point measurement.		
TOPEX TEC	-TEC measured by the TOPEX altimeter at a point.		
STDDEV	-The standard deviation caused by the 12 second averaging.		
CLM TEC	-PRISM Climatology (no GPS data) TEC value at the point.		
RTA TEC	-PRISM Real Time Adjustment (with GPS data) TEC value at the		
	point.		
Distance	-Distance to the closest GPS station from the TOPEX		
	measurement.		
Station Lat	-The closest station Latitude.		
Station Lon	-The closest station Longitude.		
Station	-The name of the closest station.		

Table 2.4: TOPEX\_Comparison6.pro Output File Variables.

# Chapter 3

### The Results

# 3.1 Local Time Analysis

In doing this analysis the error distributions were sorted by local time. In Figure 3.1 the error for the climatology and the RTA runs were plotted with respect to local time. A further subdivision was made by magnetic latitude to see the contribution of error made by each of the individual models within PRISM (LOWLAT, MIDLAT, HIGHLAT). Table 3.1 shows how the subdivisions were binned.

Model	Magnetic Latitude Range	
LOWLAT	-32 ≤ Value ≤ 32	
MIDLAT	32 < Value < 51 and -51 < Value < -32	
HIGHLAT	$Value \ge 51$ and $Value \le -51$	

Table 3.1 Magnetic Latitude Range for Comparisons

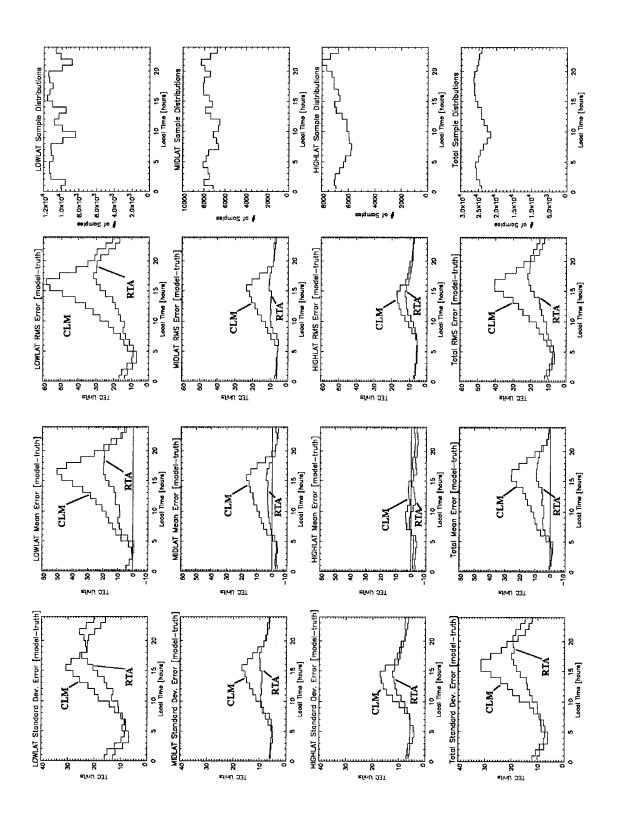


Figure 3.1: Model Error VS Local Time (CLM=Climatology, RTA=Real Time Adjustment)

The first thing to take note of is the uniformity in the sampling of local times in the distribution plots. There is a slight dip in the total number of samples around 10 LT but it is not enough to affect the results. In the error plots you can clearly see the time of day when the GPS data will really help improve the climatology. The LOWLAT model showed the largest improvement in TEC. There is an average improvement of 5.4 TEC units (RMS error) in the nighttime and a 17.4 TEC units (RMS Error) improvement during the daytime. In the MIDLAT model there is 1 TEC unit (RMS Error) of improvement during the night and an improvement of 8.7 TEC units (RMS Error) during the daytime. The HIGHLAT model showed an improvement of .9 TEC units (RMS Error) during the nighttime and an improvement of 4.2 TEC units (RMS Error) during the daytime. Overall there is a considerable improvement in all three models during the daytime when the ionospheric TEC values are the largest and there is also an improvement in the nighttime TEC when the ionospheric TEC values are the smallest. Overall, the largest improvements occurred primarily in the LOWLAT region during the daytime.

#### 3.2 Distance From Station

The next set of distributions were distance from station VS TEC error. In Figure 3.2 these distributions were plotted for the entire time period (day 35 – day 200). The data were binned in 100km sections where 0km to 100km is the first of the 70 bins.

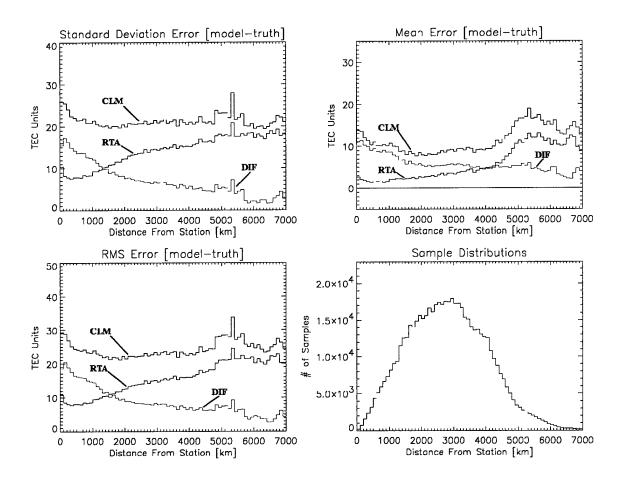


Figure 3.2: Distance From Station Verses Error (CLM=Climatology, RTA=Real Time Adjustment, DIF = CLM - RTA which is the improvement over climatology)

By looking at the sample distributions you can see that there are relatively few samples near the stations. This is primarily due to the fact that the majority of the stations are several kilometers inland. As for the error, the red histogram plot shows the improvement over climatology. This plot shows the extent to which the driver stations affect the model output as a function of distance. You can see that the GPS data overall has a positive effect and that the largest improvement happens near the station. However, there is a slight inconsistency in these plots is with the error that is very close to the stations

which can be seen in the first three bins (0km to 300km). In both the climatology (CLM) and real time adjustment (RTA) histogram plots there is a slight increase in the errors that was not expected. To determining what was going on in this region the spatial distribution of the errors larger than 20 TEC units were plotted on a global map (Figure 3.3).

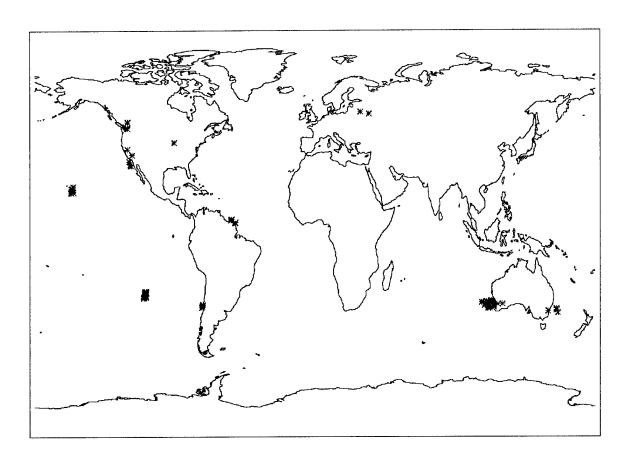


Figure 3.3: Distribution of TEC error larger than 20TEC units from 0-300km.

Figure 3.3 shows that the majority of the GPS stations did play some part in the large errors. However, there is a considerable amount of error associated with the station PERT in Australia. This station potentially does have some problems and is being looked at it more closely (Decker, personal communication). Another factor to take into consideration in the 0 to 300km region is the fact that the distance calculation is calculated from TOPEX data position to the station location and not the actual location of the GPS data given to PRISM. The actual data location can be up to 300km away from GPS receiver

for a GPS satellite elevation angel of 45 degrees. It is only when the elevation angle is 90 degrees (GPS satellite overhead) that the data location matches the station location (see figure 2.1 for the geometry).

### 3.3 Kp Distributions

The final analysis performed was to examine the errors as a function of magnetic activity. To do this the error distributions from the previous analysis (Error VS Distance From Station) were subdivided into thirteen ten day periods (see appendix D). This division permits sampling at all local times. Three hour Kp data were obtained from NOAA's National Geophysical Data Center. In figure 3.4, the three hour Kp values are plotted and are divided into thirteen plots that correspond to the thirteen data periods.

By examining these 10 day distributions and taking note of the general level of magnetic activity for that period there was a small trend. In periods where there was a small amount of magnetic activity, the errors were generally smaller than those with large amounts of activity. This can be seen by comparing the following two ten day time periods (Figures 3.5 & 3.6). The first period is of day 101-112 which has several times of high Kp and the second period is of day 161-173 which has relitivly low Kp for the entire period.

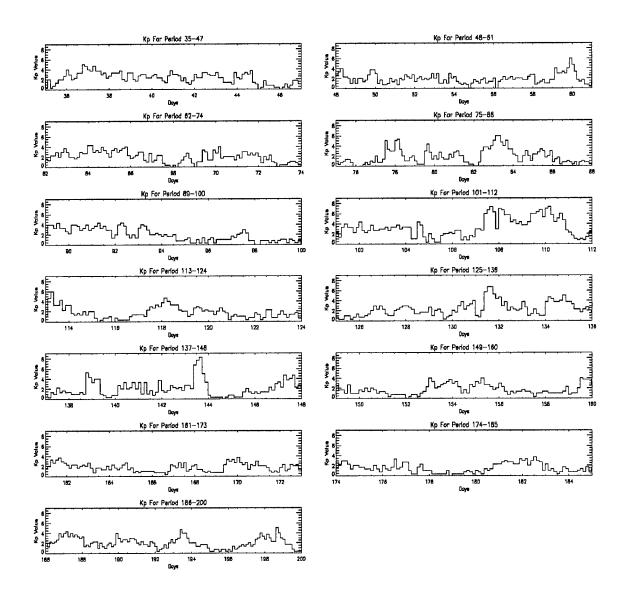


Figure 3.4: Three-Hour Kp Values.

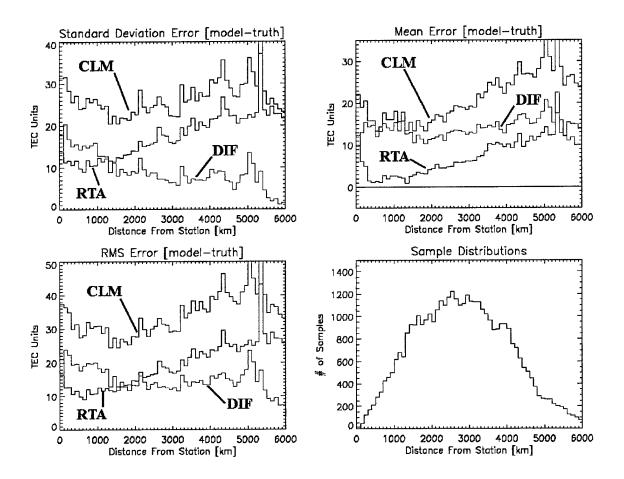


Figure 3.5: Error VS Distance From Station for period 101-112. (CLM = Climatology, RTA=Real Time Adjustment, DIF = CLM - RTA which is the improvement over climatology)

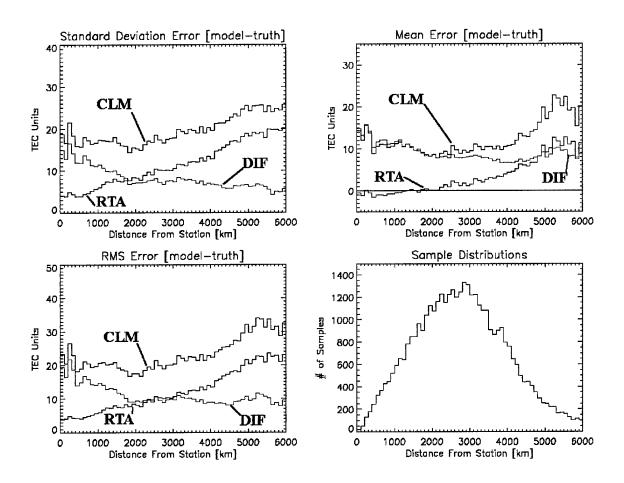


Figure 3.6: Error VS Distance From Station for period 161-173.

(CLM=Climatology, RTA=Real Time Adjustment, DIF = CLM · RTA which is the improvement over climatology)

The first thing to notice from these plots is how the climatology varied. When the Kp was high the climatology RMS error was on average 32.6 TEC units but for low Kp the average RMS error was 21.2 TEC units. Another trend was the level of error for the RTA PRISM run. For the high Kp period it was evident that GPS data did improve the climatology RMS on average by 14.3 TEC units and the low Kp period had an overall RMS improvement of 10.4 TEC units. However, even though the high Kp period showed the greatest improvement, the

RMS errors for this period were on average 18.2 TEC units where as the low Kp period error was only 10.8 TEC units. To examine the effect of Kp and its effect on the overall error a second more detailed analysis was performed.

## 3.4 Second Kp Analysis

Errors were binned by Kp instead of days, into three. The first bin is for minimum Kp with values equal to 0 to less than 3, the second bin is for moderate Kp with values from equal to 3 to less than 6, and the third for high Kp with values from equal to 6 to equal to 9. In Figure 3.7 the error distributions of the three bins are plotted with respect to local time. In appendix E these error are further broken down into the individual model components in PRISM (LOWLAT, MIDLAT, HIGHLAT).

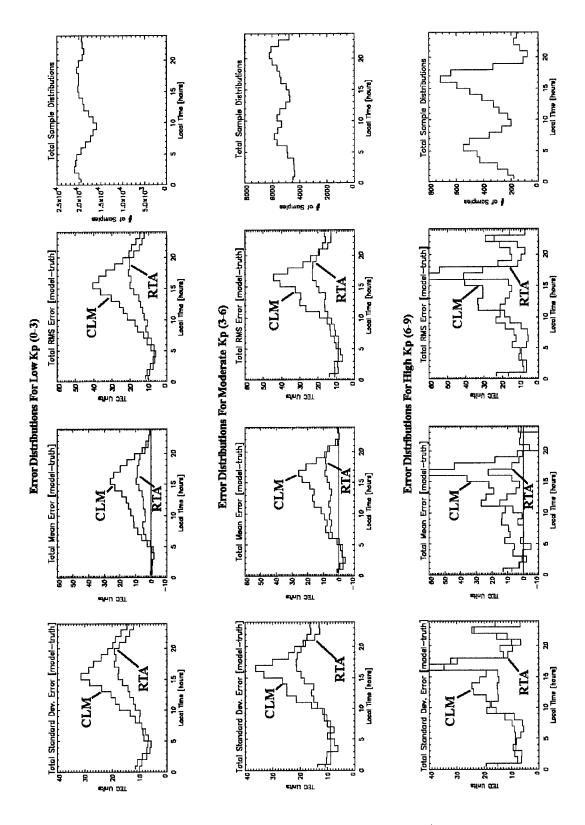


Figure 3.7: Errors by Magnetic Activity VS Local Time. (RTA=Real Time Adjustment, CLM=Climatology)

These distributions look very similar to the previous local time distributions, but there are some small differences. For instance there is a noticeable difference between the low Kp and mid Kp errors around nine local time. As Kp increases there is an increase in the level of error. Also, there is a slight increase in the overall error as the magnetic activity increases. To better see the differences Figure 3.8 puts all the RMS summary errors onto two plots, one for climatology (CLM), and the other for the real time adjustment (RTA).

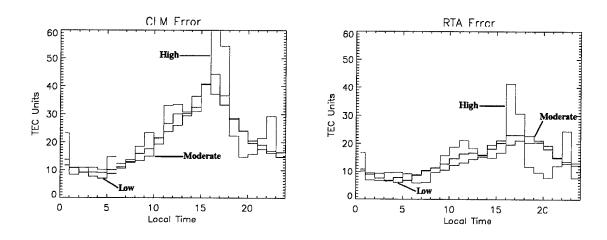


Figure 3.8: RMS Errors (Low=Low Kp, Moderate=Moderate Kp, High=High Kp)

From this figure you can clearly see that the error does increase with increasing Kp between low and moderate values. However, the high Kp is very sporadic showing improvements during some times (example: local time 18-22) and large errors during others (example: local time 16-17). This variation is probably due to the few number of samples at high Kp values. Table 3.2 below shows the number of total samples in each of the three distributions.

Kp	# of Samples	% of total samples
Low	459,758	77.5%
Moderate	125,288	21.1%
High	8,194	1.4%

Table 3.2: Kp Sample Distributions

## Chapter 4

## Summary & Conclusions

### 4.1 Summary

PRISM was developed for the Department of Defense to provide an accurate, real time specification of the ionosphere on a global scale. The climatology portion of PRISM PIM (Parameterized Ionosphere Model) requires the following indices to generate an output: Universal Time (UT), Geomagnetic Activity (Kp); Inerplanetary Magnetic Field (IMF) components (By, Bz); and the 10.7 cm solar flux (F10.7). If real time data are available, PRISM uses a Real Time Adjustment (RTA) algorithm to adjust the climatology by using weighting functions. These RTA data can be obtained from three sources: TEC data from GPS satellites, in situ measurements by DMSP, and ionospheric soundings made by the Digital Ionospheric Sounding (DISS) network. The real time input parameters include: Total Electron Content (TEC); ion drift velocities; in situ number density; ion/electron temperatures; fractional  $He^+, H^+, O^+$  content; ionospheric layer critical frequencies (foF2, foF1, and foE); and peak electron density heights (hmF2, hmF1, and hmE).

The objective of this study was to validate PRISM when driven with and without Global Positioning Satellite (GPS) TEC data. In this study PRISM was run twice (Run 1= No GPS data CLM run, Run 2=With GPS data RTA run) at hourly intervals during the period of February 4<sup>th</sup> 2002 (day 35) – July 19 2002

(day 200). This large period would provide a validation of PRISM that takes into account both the seasonal and geomagnetic variations. It should be noted that a few days and hours were removed from this period due to gaps in the GPS data but this did not adversely affect the results.

The data used for the validation is vertical TEC data from the TOPEX/Poseidon satellite. This satellite has an orbit of 1336 km and an inclination of 66 degrees. This provides worldwide (over-ocean) coverage of vertical TEC within a longitude range of 0 to 360 degree and a latitude range of -66 to 66 degrees. The vertical TEC measurements are taken every second however for this study data were averaged over 12 seconds to reduce noise in the data and reduce the number of comparisons. This 12 second data are then compared to the two PRISM runs (CLM & RTA) and the errors were analyzed.

#### 4.2 Conclusion

In this study 593,240 vertical TEC measurements were compared to the two PRISM runs for a total of 1,186,480 comparisons. The first analysis of the errors was to look at how the overall error varied with local time. In Figure 4.1, the errors are summarized for the entire period. In this figure there are three sets of distributions. The first two bars of all local time combined, the second two are of the day time local times which range from 7:00LT to 19:00LT, and the last two are of night time local times which range from 20:00LT to 6:00LT.

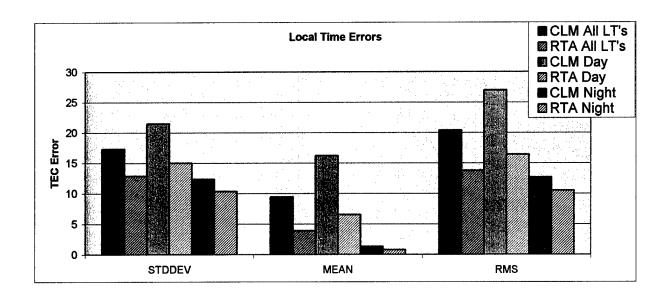
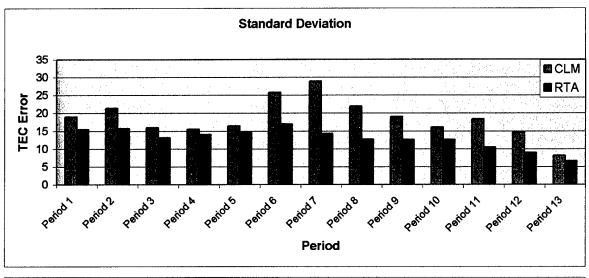
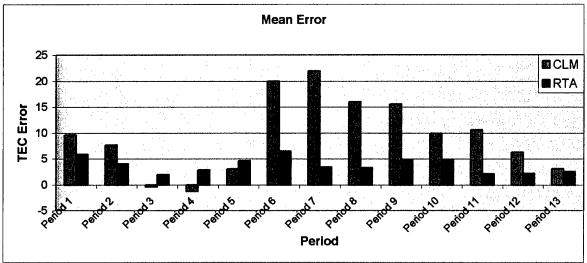


Figure 4.1: Total Error VS Local Time.

Figure 4.1 illustrates how the GPS data improve the climatology by 6.7 TECU RMS. Separating the errors by day and night you can see that the largest improvement of 10.6 TECU RMS is during the daytime when the TEC is the largest, and the smallest improvement of 2.1 TECU RMS happens during the nighttime when the TEC is the lowest.

The second error analysis evaluated the error changes with distance from the station. These errors were divided into 13 ten-day periods to help aid in seeing the seasonal and magnetic variations. Figure 4.2 summarizes the errors by these periods.





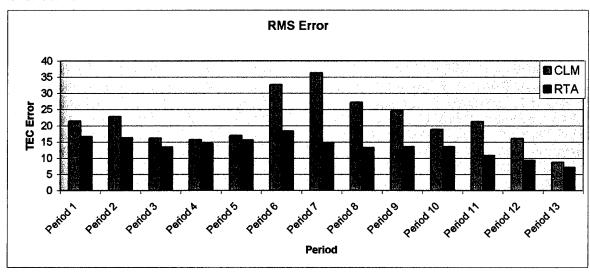


Figure 4.2: Summary Errors for the Thirteen Periods (10 days each)

From this distribution it is clear that error does vary as a function of time period or that it has a seasonal component. To explore this further a plot was made of how the daily TEC varied as a function of time. Figure 4.3 shows the daily TEC was plotted along a single meridian (zero longitude) for all latitudes and times. The TEC data are obtained by interpolating 1 degree TEC values along this meridian throughout the CLM PRISM runs.

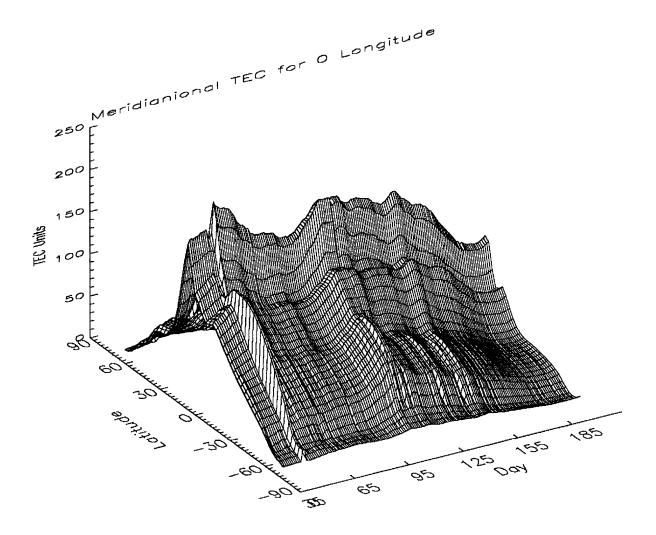


Figure 4.3: Daily TEC Along a Single Meridian (zero longitude).

When comparing Figure 4.3 with the errors in Figure 4.2, the size of the error is directly proportional to the daily maximum of TEC. So relatively speaking, the higher the daily maximum in TEC the higher the error in TEC.

The third analysis evaluates the error variation occurring with different magnetic activity. In this analysis there was a small trend as Kp increases. As the Kp ranges from low to moderate there was a slight increase in the overall error, which was on the order of a few TEC units. High Kp values perform much worse that both the moderate and low Kp the majority of times, but there were a few times where it did much better (see Fig 3.8). Since the high Kp samples only made up 1.4% of the total distribution and that the latitudes above 66 degrees were not sampled, nothing conclusive can be said about this.

Finally, the error within the three different models within PRISM (LOWLAT, MIDLAT, HIGHLAT) are summarized. The figures (Figures 4.4 & 4.5) below shows the overall errors for the model and how each of the models in PRISM contributed to the total error.

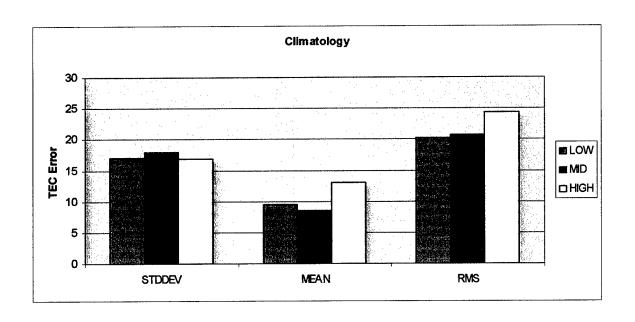


Figure 4.4: Climatology Error by PRISM Model (LOWLAT, MIDLAT, HIGHLAT)

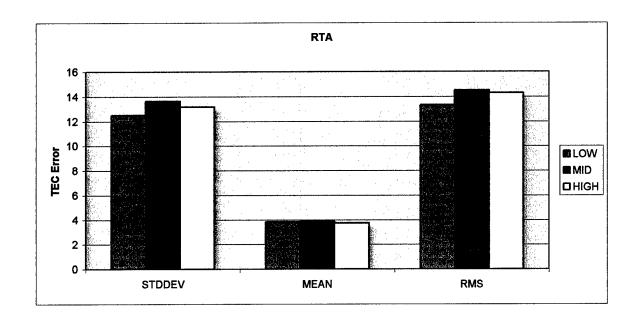
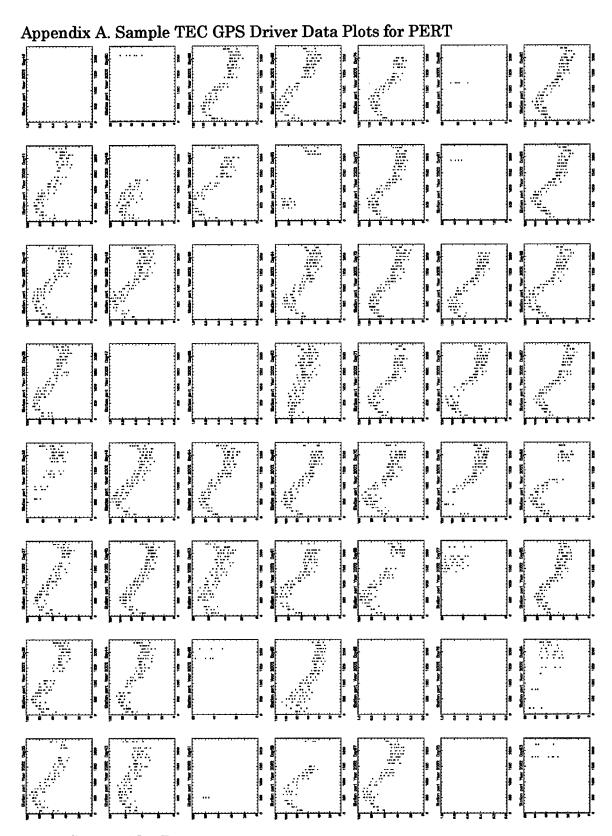


Figure 4.5: Real Time Adjustment (RTA) Error by PRISM Model (LOWLAT, MIDLAT, HIGHLAT)

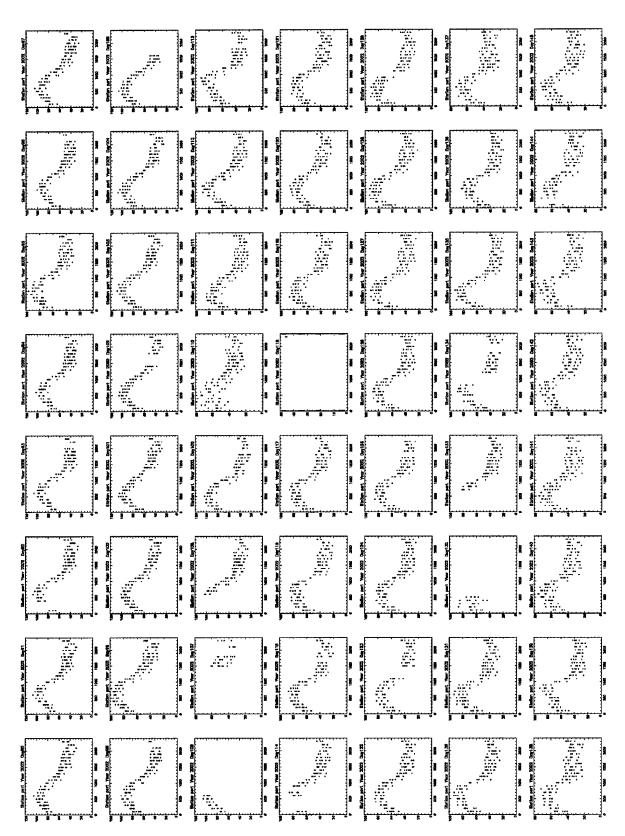
Here you can see the overall error contribution by the three different models.

These figures show that when PRISM is given driver data the error of all three models is reduced from 20 to 24 TECU to 13 to 14 TECU.

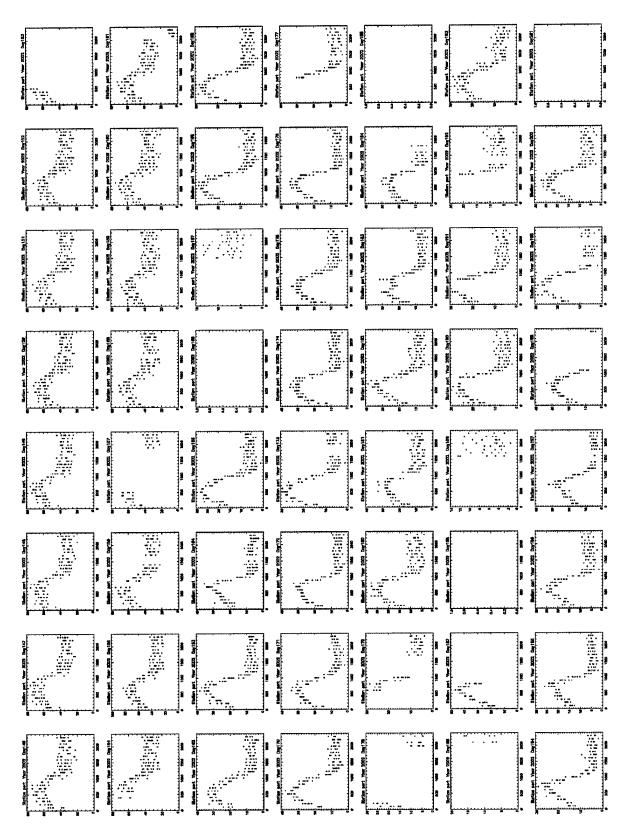
Having an accurate real time specification of the ionosphere is an essential first step to being able to forecast the ionosphere. This thesis did a thorough analysis of the real time ionosphere specification by PRISM using the same input parameters currently in use by the Air Force Weather Agency (AFWA). The intent was to make this performance analysis a useful resource to AFWA. This work also laid down the benchmark for the next generation models to be compared to. Using this validation data set and the algorithms developed in this research, furture models can be compared to PRISM to determine their improvement in accuracy. The benefits of having an accurate specification and forecast of the ionosphere would be to greatly enhance the capabilities of many DoD systems. Improvements could be made to remote sensors affected by the ionosphere which include inaccurate position readings from GPS satellites, corrections could also be made for high frequency communication disturbances, and it would be possible to predict communication outages. This knowledge would serve as a force multiplier for commanders allowing them to maximize the use of their resources and plan for possible problems caused by space weather.



PERT Station Plot Days 33-90



PERT Station Plot Days 90-145



PERT Station Plot Days 146-200

## Appendix B: Main Analysis Program

```
*************************
;IDL program that reads 24 PRISM files into a 3d array.
Modified to loop over several times.
; Modified to check station availability.
Creates a validation file containing all arrays created.
:MAIN LOOP
FOR DAYI=35,250 DO BEGIN
GET_FILE:
;**************For a standard input***********
TEC3D=FLTARR(24.46.91)
TEC3D_clm=FLTARR(24,46,91)
TEC3D_rta=FLTARR(24,46,91)
daystart=DAYI
I=daystart
FOR I=daystart,daystart DO BEGIN ;DAY loop
file_day=string(I,'(I3.3)')
file_day=strcompress(file_day,/remove_all)
FOR TYPE=0, 1 DO BEGIN ; RTA CLM Loop
FOR II=0,23 DO BEGIN ;HOUR loop
BAD hour adjustment
;;IF (II EQ 18) THEN BEGIN
;;II=II+1
;;ENDIF
file_hour=string(II,'(I2.2)')
file hour=strcompress(file hour./remove all)
IF (TYPE EQ 0) THEN BEGIN
tvpestring='CLM'
ENDIF
IF (TYPE EQ 1) THEN BEGIN
typestring='RTA'
ENDIF
FILE='E:\'+typestring+'\d'+file_day+'h'+file_hour+'\data.in'
FILE=STRCOMPRESS(FILE,/REMOVE_ALL)
GET_LUN,LUN
START:
OPENR, LUN, FILE, ERROR = err
; If err is nonzero, something happened. Print the error message to
; the standard error file (logical unit -2):
IF (err NE 0) THEN BEGIN
CLOSE.LUN
     IF (II EQ 23) THEN BEGIN
     I=I+1
     II=0
     ENDIF
     IF (II LT 23) THEN BEGIN
     II=II+1
     ENDIF
```

```
file day=string(I,'(I3.3)')
file day=strcompress(file_day,/remove_all)
file hour=string(II,'(I2.2)')
file_hour=strcompress(file_hour,/remove_all)
FILE='E:\RTA\d'+file_day+'h'+file_hour+'\data.in'
GOTO, START
ENDIF
PRINT,'Reading data...Please wait...'
FILE 1='99'
YEAR=0
DAY=0
UT=0.
F10P7=0.
KP=0.
SSN=0.
SDUM="
READF, LUN, SDUM
READF.LUN.SDUM
READF, LUN, YEAR, DAY, UT, F10P7, KP, SSN
UT=UT/3600.
; Read the primary coordinate system type, output grid type, and plasmasphere
; flag from the PIM output file
IDUM=0
READF, LUN, SDUM
READF, LUN, SDUM
READF, LUN, IDUM
CRDTYP=IDUM MOD 10
GRDTYP=(IDUM MOD 100)/10
IF(GRDTYP NE 0) THEN BEGIN
  CLOSE,LUN
 FREE LUN.LUN
 PRINT, 'File "'+FILE+'" contains the wrong type of output grid for RECCCRIT.'
 READ. 'Do you want to try another file (Y/[N])? ',SDUM
 IF STRUPCASE(STRMID(STRCOMPRESS(SDUM,/REMOVE_ALL),0,1)) EQ 'Y' THEN
BEGIN
   GOTO,GET_FILE
 ENDIF ELSE BEGIN
   GOTO, END_PULLIAM
 ENDELSE
ENDIF
; Read grid information from the PIM output file
SLAT=0.
SLON=0.
ELAT=0.
ELON=0.
NLAT=0
NLON=0
DLAT=0.
DLON=0.
READF, LUN, SDUM
READF, LUN, SDUM
READF, LUN, SLAT, ELAT, SLON, ELON, NLAT, NLON, DLAT, DLON
; Read the output data type from the PIM output file
```

```
DATTYP=0
READF.LUN.DATTYP
IF (DATTYP NE 0) AND (DATTYP NE 2) THEN BEGIN
 CLOSE,LUN
 FREE_LUN,LUN
 PRINT, File "'+FILE+" contains the wrong type of output data for RECCCRIT.'
 READ, 'Do you want to try another file (Y/[N])? ',SDUM
 IF STRUPCASE(STRMID(STRCOMPRESS(SDUM,/REMOVE_ALL),0,1)) EQ 'Y' THEN
BEGIN
  GOTO,GET_FILE
 ENDIF ELSE BEGIN
  GOTO, END_PULLIAM
 ENDELSE
ENDIF
; Read critical frequencies and heights and TEC from the PIM output file
IF DATTYP EQ 2 THEN BEGIN
 NALT=0
 READF.LUN.NALT.FORMAT='(29X,I6)'
 READF, LUN, SDUM
 ALT=FLTARR(NALT)
 READF, LUN, ALT
 EDP=FLTARR(NALT)
ENDIF
GLAT=FLTARR(NLAT.NLON)
GLON=FLTARR(NLAT,NLON)
MLAT=FLTARR(NLAT,NLON)
MLON=FLTARR(NLAT,NLON)
MLT=FLTARR(NLAT,NLON)
FOF2=FLTARR(NLAT, NLON)
HMF2=FLTARR(NLAT.NLON)
FOF1=FLTARR(NLAT, NLON)
HMF1=FLTARR(NLAT,NLON)
FOE=FLTARR(NLAT,NLON)
HME=FLTARR(NLAT, NLON)
TEC=FLTARR(NLAT,NLON)
TEC2=FLTARR(NLON,NLAT)
DUM1=FLTARR(5)
DUM2=FLTARR(7)
FOR ILAT=0,NLAT-1 DO BEGIN
 FOR ILON=0.NLON-1 DO BEGIN
  READF, LUN, SDUM
  READF,LUN,DUM1
  GLAT(ILAT,ILON)=DUM1(0)
  GLON(ILAT.ILON)=DUM1(1)
  MLAT(ILAT,ILON)=DUM1(2)
  MLON(ILAT,ILON)=DUM1(3)
  MLT(ILAT,ILON)=DUM1(4)
  IF DATTYP EQ 2 THEN BEGIN
    READF, LUN, SDUM
    READF, LUN, EDP
  ENDIF
  READF, LUN, SDUM
  READF, LUN, DUM2
```

```
FOF2(ILAT,ILON)=DUM2(0)
   HMF2(ILAT.ILON)=DUM2(1)
   FOF1(ILAT,ILON)=DUM2(2)
   HMF1(ILAT.ILON)=DUM2(3)
   FOE(ILAT,ILON)=DUM2(4)
   HME(ILAT,ILON)=DUM2(5)
   TEC(ILAT,ILON)=DUM2(6)
   TEC2(ILON,ILAT)=DUM2(6)
   TEC3D(II,ILAT,ILON)=DUM2(6)
   IF (ILON EQ 0) THEN BEGIN
   TEC3D(II,ILAT,90)=DUM2(6)
   ENDIF
 ENDFOR
ENDFOR
CLOSE,LUN
FREE_LUN,LUN
END_PULLIAM:
PRINT, 'FINISHED'
ENDFOR ;Hour loop
IF (TYPE EQ 0) THEN BEGIN
TEC3D_clm=TEC3D
ENDIF
IF (TYPE EQ 1) THEN BEGIN
TEC3D rta=TEC3D
ENDIF
ENDFOR ;TYPE Loop RTA CLM
daystart=I
station 1="
time=fltarr(10000)
latitude=fltarr(10000)
longitude=fltarr(10000)
vtec=fltarr(10000)
station=strarr(10000)
slat=fltarr(10000)
slon=fltarr(10000)
day=fltarr(10000)
f=01
file_day=string(daystart, '(13.3)')
file day=strcompress(file_day,/remove_all)
FOR II=0,23 DO BEGIN ;HOUR loop
file hour=string(II,'(I2.2)')
file hour=strcompress(file_hour,/remove_all)
FILE='E:\PRISM\no_gt_2002\rta\d'+file_day+'\d'+file_day+'h'+file_hour+'.dat'
close.1
START1:
openr,1,FILE,ERROR = err ; MUST PUT ERROR SKIP ROUTINE HERE
IF (err NE 0) THEN BEGIN
CLOSE.1
      IF (II EQ 23) THEN BEGIN
      I=I+1
      II=0
      ENDIF
```

```
IF (II LT 23) THEN BEGIN
       II=II+1
       ENDIF
       Incase there is a missing file
       day1=FIX(FILE_DAY)
       vtec(F)=0
       day(F)=day1
       latitude(F)=0
       longitude(F)=0
       TIME(F)=0
       station1=strcompress(station1,/remove_all)
       station(F)=0
       slat(F)=0
       slon(F)=0
       F=F+1
file_day=string(I,'(I3.3)')
file_day=strcompress(file_day,/remove_all)
file_hour=string(II,'(I2.2)')
file hour=strcompress(file_hour,/remove_all)
FILE='E:\PRISM\no_gt_2002\rta\d'+file_day+'\d'+file_day+'h'+file_hour+'.dat'
GOTO, START1
ENDIF
WHILE NOT EOF(1) DO BEGIN
day1=FIX(FILE_DAY)
readf,1,format='(5F9.3,A6,3f9.3)',dum,time1,lat1,lon1,vtec1,station1,dum,slat1,slon1
       vtec(F)=vtec1
       day(F)=day1
       latitude(F)=lat1
       longitude(F)=lon1
       TIME(F)=time1
       station1=strcompress(station1,/remove_all)
       station(F)=station1
       slat(F)=slat1
       IF (slon1 GT 180) THEN BEGIN
       newslon1=slon1-180
       lon1=-180+newslon1
       ENDIF
       slon(F)=lon 1
       F=F+1
ENDWHILE
ENDFOR
G=F-1
latitude=latitude(0:G)
longitude=longitude(0:G)
vtec=vtec(0:G)
time=time(0:G)
station=station(0:G)
slat=slat(0:G)
slon=slon(0:G)
dav=dav(0:G)
Plotting instructions
```

```
stationloop=['aoa1','ared','cic1','cro1','drao','eisl','gode','gol2','gold','guam','hrao','jplm','kiru','kour','k
okb', 'madr', 'mad2', 'madr', 'mcm4', 'mdo1', 'mkea', 'nlib', 'nya2', 'pert', 'pie1', 'pots', 'quin', 'sant', 'suth', 'tid
b'.'usud'.'zwen']
StationAvail=fltarr(n_elements(stationloop),24)
StationAvail_lat=fltarr(n_elements(stationloop),24)
StationAvail_lon=fltarr(n_elements(stationloop).24)
endofarray=n elements(time)-1
FOR I=0, endofarray DO BEGIN
el1=time(I)
el1=el1/100
el2=station(I)
el3=slat(I)
el4=slon(I)
       FOR II=0. N ELEMENTS(stationloop)-1 DO BEGIN
       IF (el2 EQ stationloop(II)) THEN BEGIN
       el1=fix(el1)
       StationAvail(II,el1)=[1]
       StationAvail_lat(II.el1)=[el3]
       StationAvail_lon(II,el1)=[el4]
       ENDIF
       ENDFOR
ENDFOR ; End of Station Availiability
This program reads in one 12 second data file into arrays.
dum="
Tday=fltarr(9000)
Tut=fltarr(9000)
Tlat=fltarr(9000)
Tlon=fltarr(9000)
Tlonconv=fltarr(9000)
Ttec=fltarr(9000)
Tloct=fltarr(9000)
TSTDDEV=fltarr(9000)
counter=01
dayindex=FIX(daystart)
davindex=string(davindex,'(13.3)')
dayindex=strcompress(dayindex,/remove_all)
file='E:\TOPEX Data\Processed_12_data\TOPEX_12s_'+dayindex+'.txt'
close,1
openr, 1, file, ERROR = err
       IF (err NE 0) then begin
       print.'Could not find the TOPEX data FILE'
       print,file
       ENDIF
       readf,1,dum
       WHILE NOT EOF(1) DO BEGIN
       readf,1,vear1,day1,ut1,lt1,lat1,lon1,tec1,stddev1
       Tdav(counter)=[dav1]
       Tut(counter)=[ut1]
       Tloct(counter)=[lt1]
       Tlat(counter)=[lat1]
       Tlon(counter)=[lon1]
```

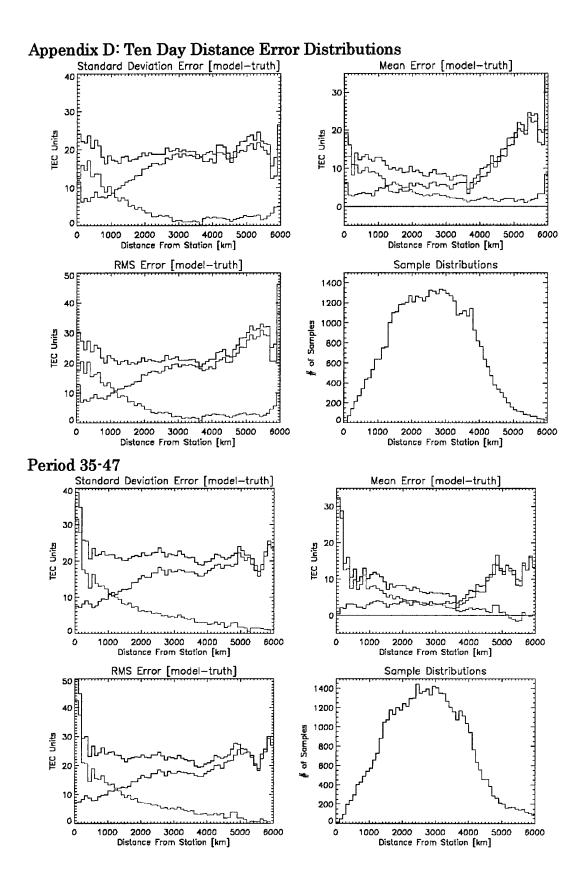
```
IF (lon1 GT 180) THEN BEGIN
      newslon1=lon1-180
      lon1=-180+newslon1
      ENDIF
      Tlonconv(counter)=[lon1]
      Ttec(counter)=[tec1]
      TSTDDEV(counter)=[stddev1]
      counter=counter+1
      ENDWHILE
Tday=Tday(0:counter)
Tut=Tut(0:counter)
Tloct=Tloct(0:counter)
Tlat=Tlat(0:counter)
Tlon=Tlon(0:counter)
Tlonconv=Tlonconv(0:counter)
Ttec=Ttec(0:counter)
TSTDDEV=TSTDDEV(0:counter)
This part fixes the removal of lon 360 from the PRISM output
;The array is FIXEDTEC(latitude=46, longitude=91)
;latitude array(0)=-90 array(45)=90
;longitude array(0)=0 array(90)=360
PRISMTECARR_clm=fltarr(counter+1)
PRISMTECARR_rta=fltarr(counter+1)
MAGLAT=fltarr(counter+1)
MAGLON=fltarr(counter+1)
interlat=(tlat+90)/4
interlon=tlon/4
      FOR III=0,counter DO BEGIN
      PRISMTEC1=interpolate(TEC3D_clm,[Tut(III)],[interlat(III)],[interlon(III)],/grid)
      PRISMTEC2=interpolate(TEC3D_rta,[Tut(III)],[interlat(III)],[interlon(III)],/grid)
      MAGLAT1=interpolate(MLAT,[interlat(III)],[interlon(III)],/grid)
      MAGLON1=interpolate(MLON,[interlat(III)],[interlon(III)],/grid)
      PRISMTECARR_clm(III)=PRISMTEC1
      PRISMTECARR rta(III)=PRISMTEC2
      MAGLAT(III)=MAGLAT1
      MAGLON(III)=MAGLON1
      ENDFOR
END of Interpolation
;Calculates the distance to the nearest station.
distancearray=fltarr(n_elements(stationloop))
dis=fltarr(n elements(ttec)) ;contains distances
disstationlat=fltarr(n_elements(ttec))
disstationlon=fltarr(n_elements(ttec))
FOR K=0, counter DO BEGIN sloop over every data point
discounter=n_elements(stationloop)-1
Tdislat=Tlat(K)
Tdislon=Tlonconv(K)
time=Tut(K)
timeround=round(time)
IF (timeround EQ 24) THEN BEGIN
```

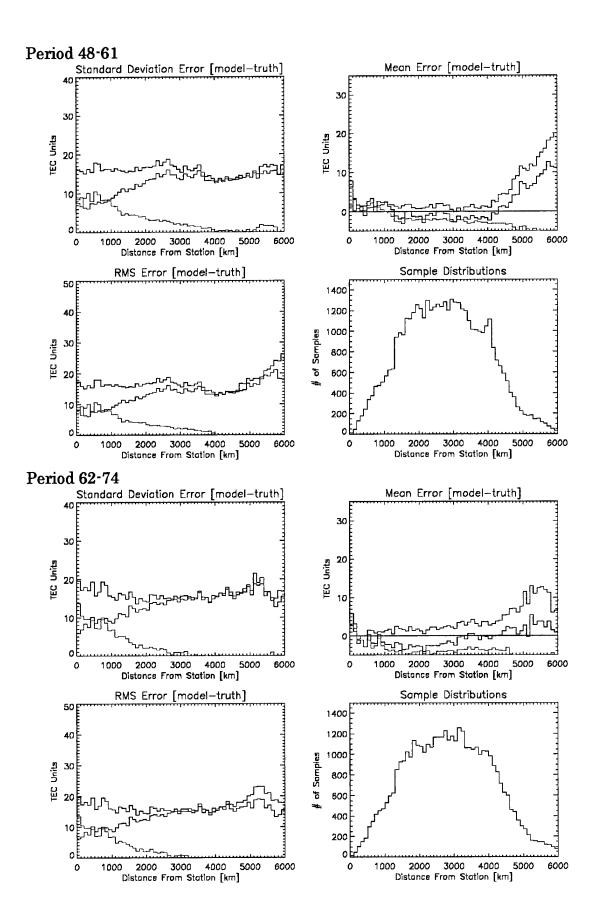
```
timeround=23
ENDIF
      FOR KK=0, discounter DO BEGIN ;Sets array to 20000km
      distancearray(KK)=20000
      ENDFOR
             FOR KKK=0, discounter DO BEGIN
             test=StationAvail(KKK,timeround)
                   IF (test EQ 1) THEN BEGIN
      dis1=MAP_2POINTS(StationAvail_lon(KKK,timeround),StationAvail_lat(KKK,timeroun
d), Tdislon, Tdislat,/meters)
                   dis1 = dis1/1000
                   distancearray(KKK)=dis1
                   ENDIF
             ENDFOR
dis(K)=min(distancearray)
disstation1=STATIONLOOP(WHERE(DISTANCEARRAY EQ MIN(DISTANCEARRAY)))
disstation(K)=disstation1(0)
disstationlat1=slat(where(station EQ disstation1(0)))
disstationlat(K)=disstationlat1(0)
disstationlon1=slon(where(station EQ disstation1(0)))
disstationlon(K)=disstationlon1(0)
ENDFOR
close.2
      writefile='ValidationDay1a_'+dayindex+'.txt'
      openw,2,writefile
      printf,2,'DAY UT
                      LT
                             LAT
                                  LON
                                          MLat MLon TTEC STDDEV P_CLM
              DSLat DSLon DStation'
P RTA DIST
             counter=counter-2
                   FOR W=0,counter DO BEGIN
                   days=day1
                          a=strcompress(days,/remove_all)
                   uts=tut(W)
                          b=strcompress(uts,/remove_all)
                   tlocts=tloct(W)
                          c=strcompress(tlocts,/remove_all)
                   tlats=tlat(W)
                          d=strcompress(tlats,/remove_all)
                   tlons=tlon(W)
                          e=strcompress(tlons,/remove_all)
                   maglats=maglat(W)
                          f=strcompress(maglats,/remove_all)
                   maglons=maglon(W)
                          g=strcompress(maglons,/remove_all)
                   ttecs=ttec(W)
                          h=strcompress(ttecs,/remove_all)
                   tstddevs=tstddev(W)
                          i=strcompress(tstddevs,/remove_all)
                   prism_clms=prismtecarr_clm(W)
                          j=strcompress(prism_clms,/remove_all)
                   prism_rtas=prismtecarr_rta(W)
```

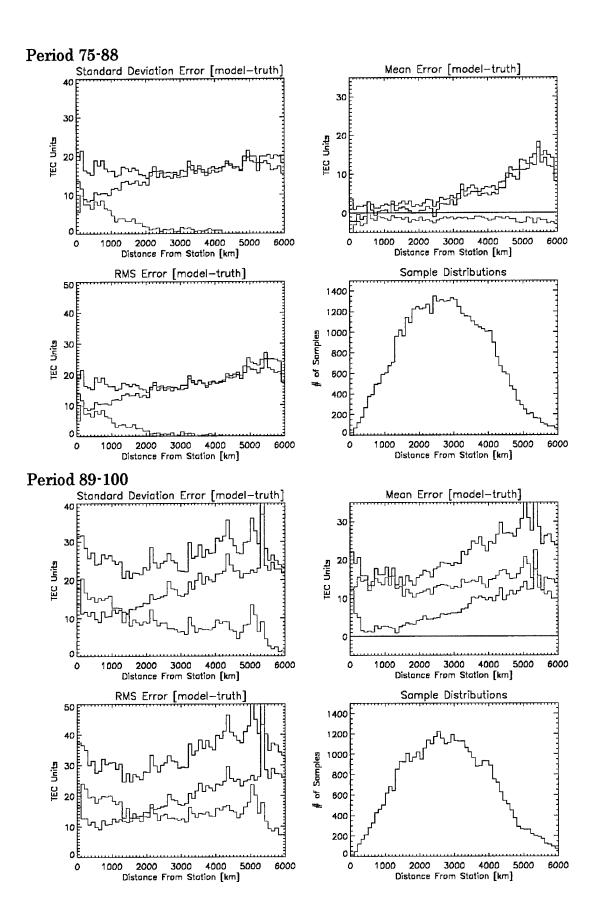
ENDFOR END

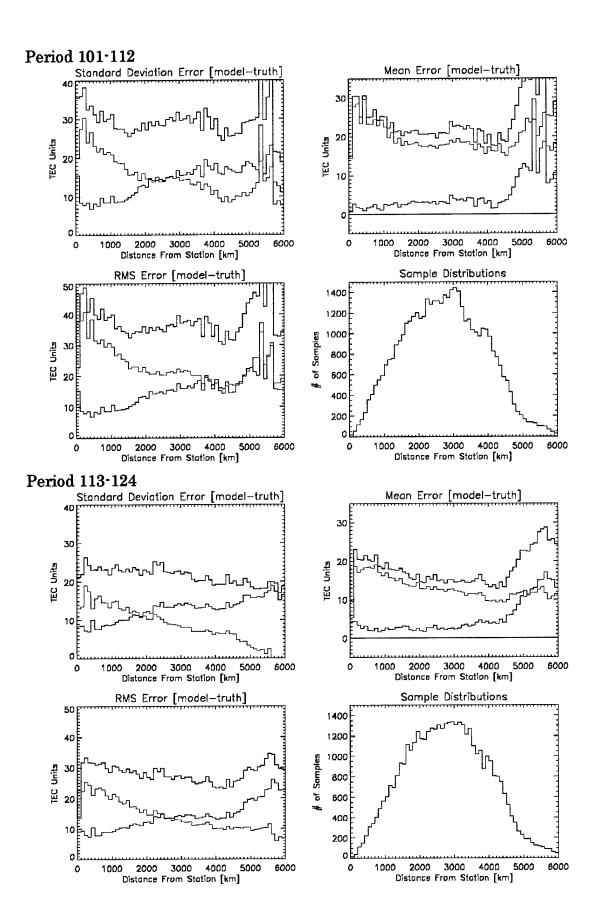
## Appendix C: Sample Validation File

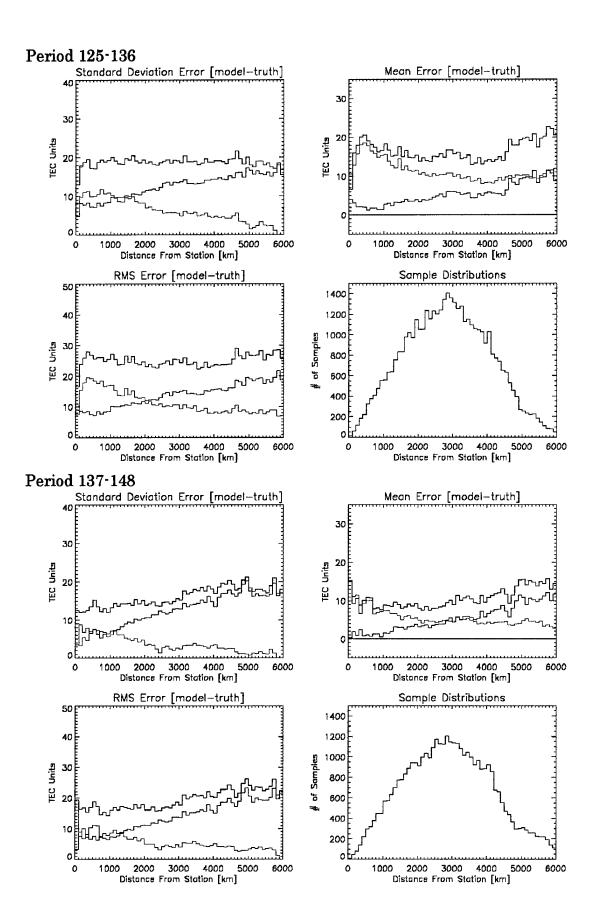
```
DAY UT LT LAT LON MLat MLon TTEC STDDEV P_CLM P_RTA DIST DSLat DSLon DStation
35 1.15324 2.34969 -66.1400 17.9468 -60.6322 59.8250 29.9167 2.87107 22.6815 25.7079 3332.87 -77.8400 187.450 mcm4
35\ 1.15685\ 2.46357\ \cdot 66.1200\ 19.6007\ \cdot 60.9012\ 60.9984\ 29.1667\ 2.03443\ 22.6667\ 25.7810\ 3304.89\ \cdot 77.8400\ 187.450\ \mathbf{mcm4}
35 1.16047 2.57718 -66.0811 21.2506 -61.1592 62.1903 27.0000 2.08167 22.6152 25.8540 3278.36 -77.8400 187.450 mcm4
35 1.16409 2.69036 -66.0235 22.8941 -61.4055 63.4004 28.8333 2.79384 22.5459 25.9354 3253.30 -77.8400 187.450 mcm4
35\ 1.16771\ 2.80295\ -65.9473\ 24.5287\ \cdot 61.6383\ 64.6204\ 27.8333\ 2.15381\ 22.4005\ 25.9035\ 3229.77\ \cdot 77.8400\ 187.450\ mcm4
35 1.17148 2.91942 ·65.8481 26.2191 ·61.8661 65.9028 27.1667 2.03443 22.1188 25.6283 3206.93 ·77.8400 187.450 mcm4
35\ 1.17570\ 3.04871\ \cdot 65.7139\ 28.0952\ \cdot 62.1050\ 67.3575\ 27.1667\ 2.22985\ 21.7979\ 25.3188\ 3183.35\ \cdot 77.8400\ 187.450\ mcm4
35 1.17974 3.17123 ·65.5622 29.8723 ·62.3239 68.7693 26.8333 2.26691 21.5351 25.0788 3162.85 ·77.8400 187.450 mcm4
35 1.18369 3.28928 -65.3938 31.5839 -62.5207 70.1620 24.7500 1.63936 21.2764 24.8462 3144.76 -77.8400 187.450 mcm4
35 1 18733 3 39686 ·65.2195 33.1430 ·62.6913 71.4626 24.4167 2.09993 21.0706 24.6660 3129.90 ·77.8400 187.450 mcm4
35 1.19248 3.54638 -64.9439 35.3086 -62.9079 73.3230 26.0000 2.79881 20.7985 24.4368 3111.86 -77.8400 187.450 mcm4
35\ 1.19751\ 3.68912\ \cdot 64.6429\ 37.3742\ \cdot 63.0881\ 75.1620\ 27.2500\ 1.96320\ 20.5162\ 24.1452\ 3097.68\ \cdot 77.8400\ 187.450\ mcm4
35 1.20640 3.93339 ·64.0351 40.9048 ·63.3485 78.4940 27.3333 3.09121 20.0119 23.5669 3081.13 ·77.8400 187.450 mcm4
35 1.21245 4.09395 ·63.5751 43.2225 ·63.5094 80.8470 29.5833 2.95687 19.7058 23.2624 3075.35 ·77.8400 187.450 mcm4
35 1,21695 4,20956 -63.2065 44.8893 -63.5978 82.6155 26.5000 3.64005 19.4988 23.0633 3074.53 -77.8400 187.450 mcm4
35 1.22124 4.31709 ·62.8342 46.4377 ·63.6580 84.3209 27.9167 2.84190 19.3205 22.9007 3076.48 ·77.8400 187.450 mcm4
35 1.22528 4.41593 ·62.4681 47.8597 ·63.6951 85.9408 30.8333 3.18416 19.1528 22.7484 3080.49 ·77.8400 187.450 mcm4
35 1.22890 4.50220 ·62.1273 49.0995 ·63.7224 87.4172 30.0000 2.85774 19.0193 22.6233 3086.02 ·77.8400 187.450 mcm4
35 1.23251 4.58651 -61.7745 50.3100 -63.7346 88.8496 28.5000 2.17945 18.8723 22.4958 3093.34 -77.8400 187.450 mcm4
35 1.23613 4.66888 ·61.4105 51.4913 ·63.7303 90.2520 29.3333 2.89636 18.7243 22.3472 3102.40 ·77.8400 187.450 mcm4
35\ 1.23975\ 4.74933\ \cdot 61.0355\ 52.6437\ \cdot 63.7155\ 91.6658\ 29.9167\ 2.17786\ 18.5939\ 22.2137\ 3113.21\ \cdot 77.8400\ 187.450\ mcm4
35 1.24336 4.82786 -60.6500 53.7675 -63.6882 93.0896 29.5833 4.27119 18.4698 22.1005 3125.75 -77.8400 187.450 mcm4
35 1.24697 4.90451 -60.2543 54.8631 -63.6441 94.5149 29.9167 3.98870 18.3437 21.9833 3140.00 -77.8400 187.450 mcm4
35\ 1.25059\ 4.97930\ -59.8490\ 55.9308\ -63.5835\ 95.9413\ 30.8333\ 2.70288\ 18.2160\ 21.8627\ 3155.92\ -77.8400\ 187.450\ mcm4
35\ 1.25420\ 5.05228\ \cdot 59.4344\ 56.9712\ \cdot 63.5069\ 97.3690\ 30.9167\ 2.49861\ 18.1277\ 21.7787\ 3173.50\ \cdot 77.8400\ 187.450\ mcm4
35 1.25781 5.12347 ·59.0108 57.9848 ·63.4145 98.7977 32.8333 3.67045 18.0430 21.6976 3192.71 ·77.8400 187.450 mcm4
35 1.26143 5.19290 -58.5787 58.9721 -63.3068 100.227 30.9167 3.59301 17.9594 21.6170 3213.51 -77.8400 187.450 mcm4
35 1,26504 5,26063 -58,1383 59,9338 -63,1842 101,658 31,3333 2,56038 17,8770 21,5369 3235,88 -77,8400 187,450 mcm4
35 1.26865 5.32668 -57.6901 60.8704 -63.0655 103.104 32.3333 2.68742 17.7936 21.5022 3259.78 -77.8400 187.450 mcm4
35 1.27226 5.39111 ·57.2343 61.7826 ·62.9302 104.519 30.2500 4.58485 17.7049 21.4749 3285.18 ·77.8400 187.450 mcm4
35\ 1.27588\ 5.45394\ -56.7713\ 62.6710\ -62.7804\ 105.924\ 29.7500\ 2.24072\ 17.6219\ 21.4460\ 3312.04\ -77.8400\ 187.450\ mcm4
35 1.27949 5.51523 ·56.3014 63.5362 ·62.6164 107.321 27.5833 2.53174 17.5447 21.4152 3340.33 ·77.8400 187.450 mcm4
35 1,28310 5,57503 ·55.8248 64.3789 ·62.4387 108.710 27.4167 3.20048 17.4768 21.3665 3370.02 ·77.8400 187.450 mcm4
35 1.28671 5.63336 ·55.3418 65.1997 ·62.2476 110.090 26.3333 4.08928 17.4182 21.2861 3401.06 ·77.8400 187.450 mcm4
35 1.29032 5.69027 ·54.8528 65.9993 ·62.0437 111.462 26.2500 2.20322 17.3643 21.1918 3433.41 ·77.8400 187.450 mcm4
35 1.29393 5.74582 ·54.3579 66.7783 ·61.8273 112.825 27.5000 3.17543 17.3150 21.0861 3467.04 ·77.8400 187.450 mcm4
35 1.29754 5.80003 ·53.8574 67.5374 ·61.5925 114.155 27.6667 3.54338 17.2667 20.9785 3501.92 ·77.8400 187.450 mcm4
35 1.30114 5.85295 -53.3516 68.2770 -61.3293 115.422 24.7500 3.63146 17.2082 20.9136 3538.00 -77.8400 187.450 mcm4
35 1,30475 5,90462 ·52.8406 68.9980 ·61.0547 116.704 25.2500 3.16557 17.1499 20.8821 3575.24 ·77.8400 187.450 mcm4
35 1.30836 5.95508 ·52.3247 69.7008 ·60.7689 117.974 24.6667 2.13437 17.0913 20.8390 3613.62 ·77.8400 187.450 mcm4
35 1.31197 6.00437 ·51.8041 70.3860 ·60.4723 119.231 24.3333 3.51979 17.0324 20.7846 3653.09 ·77.8400 187.450 mcm4
35\ 1.31557\ 6.05253\ -51.2790\ 71.0543\ -60.1653\ 120.476\ 26.6667\ 2.05480\ 16.9733\ 20.7188\ 3693.61\ -77.8400\ 187.450\ mcm4
35 1.31918 6.09959 ·50.7495 71.7062 ·59.8481 121.710 25.0833 3.32812 16.9143 20.6420 3735.17 ·77.8400 187.450 mcm4
35 1.32279 6.14560 ·50.2159 72.3422 ·59.5244 122.953 25.5000 1.97906 16.8675 20.5643 3777.70 ·77.8400 187.450 mcm4
35\ 1.32639\ 6.19058\ \cdot 49.6782\ 72.9629\ \cdot 59.1903\ 124.149\ 25.5833\ 3.22641\ 16.9366\ 20.6098\ 3821.20\ \cdot 77.8400\ 187.450\ mcm4
35 1.33000 6.23458 ·49.1368 73.5687 ·58.8428 125.294 26.0833 3.27766 17.0772 20.7389 3865.61 ·77.8400 187.450 mcm4
35 1.33360 6.27761 ·48.5917 74.1602 ·58.4846 126.425 26.6667 3.34996 17.2184 20.8695 3817.57 ·31.8000 115.720 pert
35 1,33721 6,31973 -48,0431 74,7378 -58,1162 127,541 27,0833 1,38193 17,3605 21,0021 3761,06 -31,8000 115,720 pert
35 1.34081 6.36095 ·47.4910 75.3021 ·57.7377 128.644 26.5833 2.72208 17.5037 21.1365 3705.17 ·31.8000 115.720 pert
35 1.34441 6.40131 ·46.9358 75.8535 ·57.3497 129.733 27.6667 3.06413 17.6482 21.2730 3649.94 ·31.8000 115.720 pert
35 1.34801 6.44084 ·46.3774 76.3924 ·56.9522 130.809 29.5833 3.14797 17.8024 21.4223 3595.40 ·31.8000 115.720 pert
35 1.35162 6.47957 ·45.8160 76.9192 ·56.5397 131.849 29.4167 2.56445 18.0193 21.6433 3541.59 ·31.8000 115.720 pert
35 1,35522 6.51751 -45.2517 77,4344 -56.1069 132.829 28.0000 1.35401 18.3719 22.0180 3488.53 -31.8000 115.720 pert
35 1.35882 6.55471 ·44.6847 77.9383 ·55.6670 133.798 28.4167 2.36144 18.7505 22.4233 3436.26 ·31.8000 115.720 pert
35 1.36242 6.59118 ·44.1151 78.4314 ·55.2203 134.754 29.7500 3.26917 19.1547 22.8588 3384.82 ·31.8000 115.720 pert
35 1,36602 6.62696 -43.5428 78.9140 -54.7667 135.700 29.7500 3.00347 19.5843 23.3241 3334.23 -31.8000 115.720 pert
35 1.36962 6.66206 -42.9682 79.3865 -54.3067 136.634 28.2500 3.11247 20.0387 23.8187 3284.55 -31.8000 115.720 pert
35 1.37322 6.69650 -42.3912 79.8492 -53.8405 137.557 29.0833 3.20048 20.5178 24.3422 3235.81 -31.8000 115.720 pert
35 1.37681 6.73031 ·41.8119 80.3025 ·53.3497 138.481 31.2500 2.12623 21.0488 24.9285 3188.05 ·31.8000 115.720 pert
35 1.38041 6.76352 -41.2304 80.7466 -52.8303 139.368 29.7500 2.34965 21.6123 25.5503 3141.33 -31.8000 115.720 pert
35 1.38401 6.79614 ·40.6468 81.1819 ·52.3066 140.242 31.5833 3.22641 22.1882 26.1863 3095.68 ·31.8000 115.720 pert
35 1.38776 6.82951 ·40.0368 81.6262 ·51.7568 141.138 31.0000 2.94392 22.8010 26.8632 3049.33 ·31.8000 115.720 pert
```

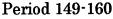


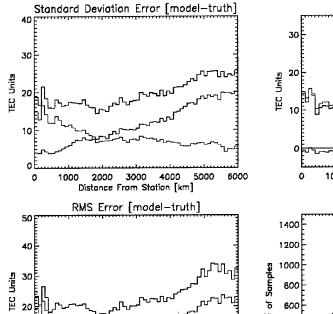


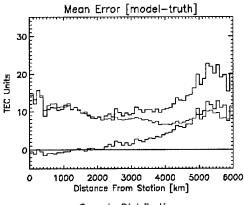


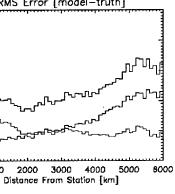


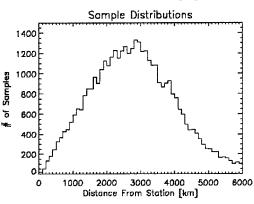










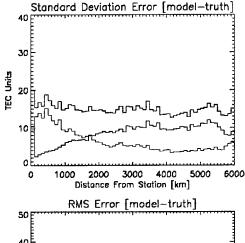


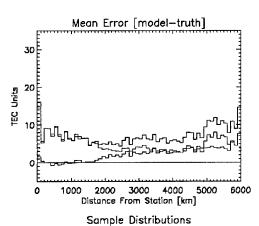
Period 161-173

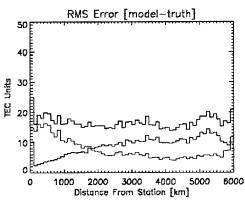
1000

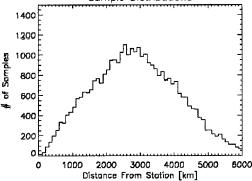
10

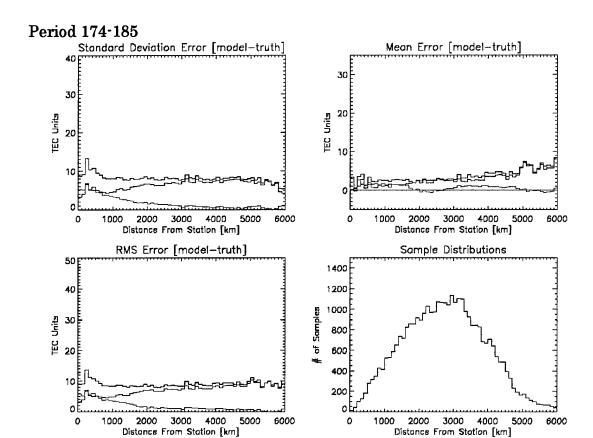
οĒ





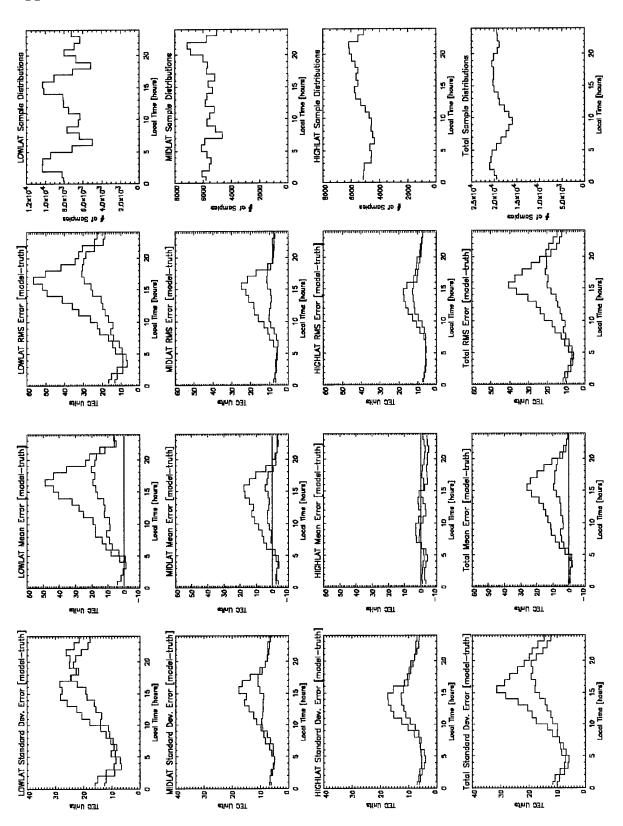


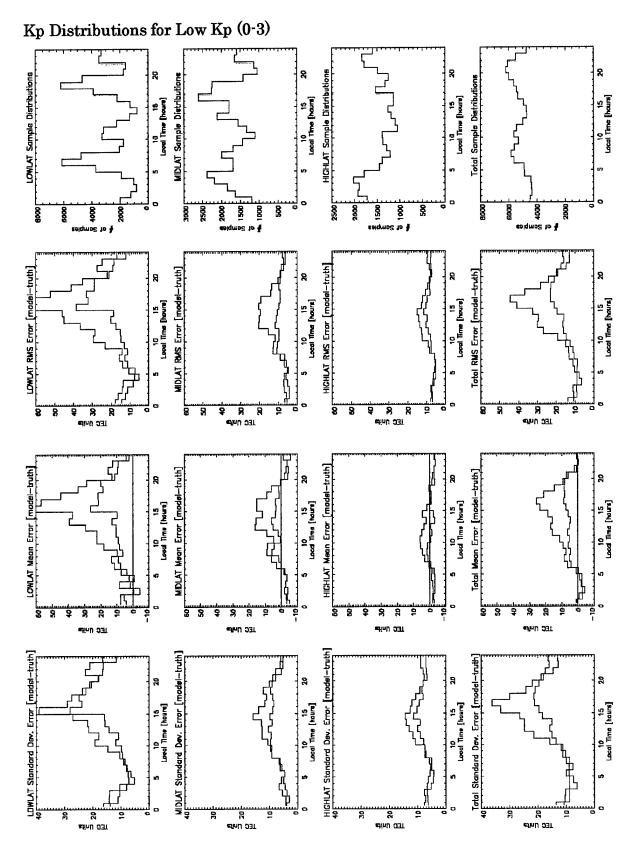




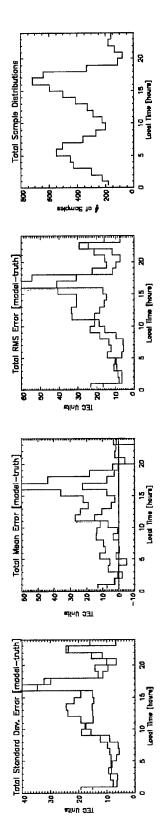
Period 186-200

Appendix E: Local Time Error Distributions by Kp





Kp Distributions for Moderate Kp (3-6)

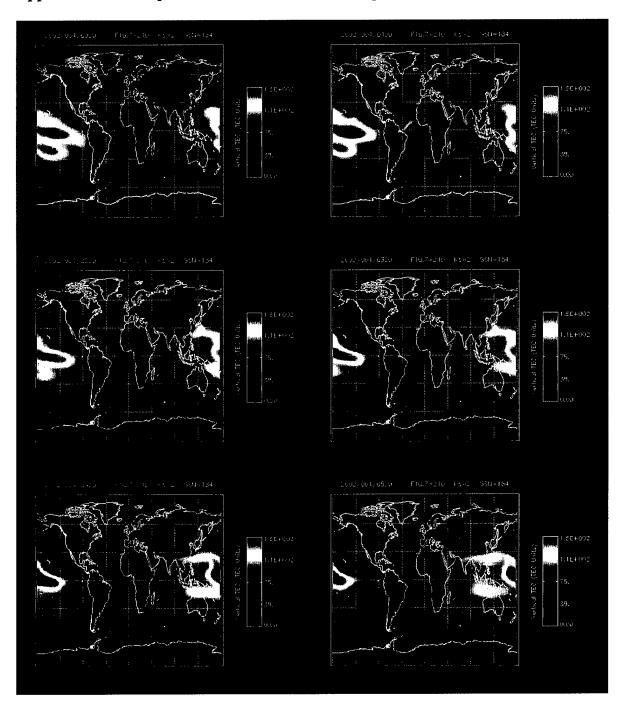


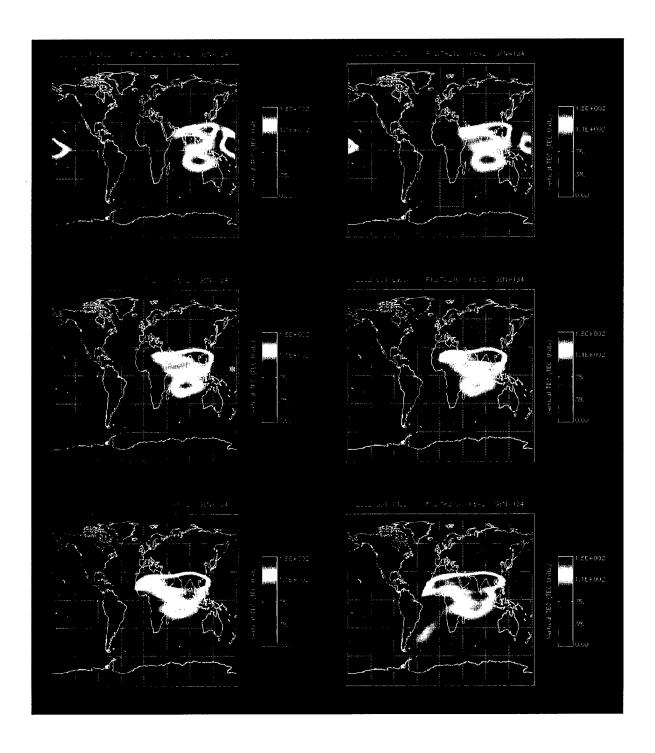
Kp Distributions for Moderate Kp (6-9)

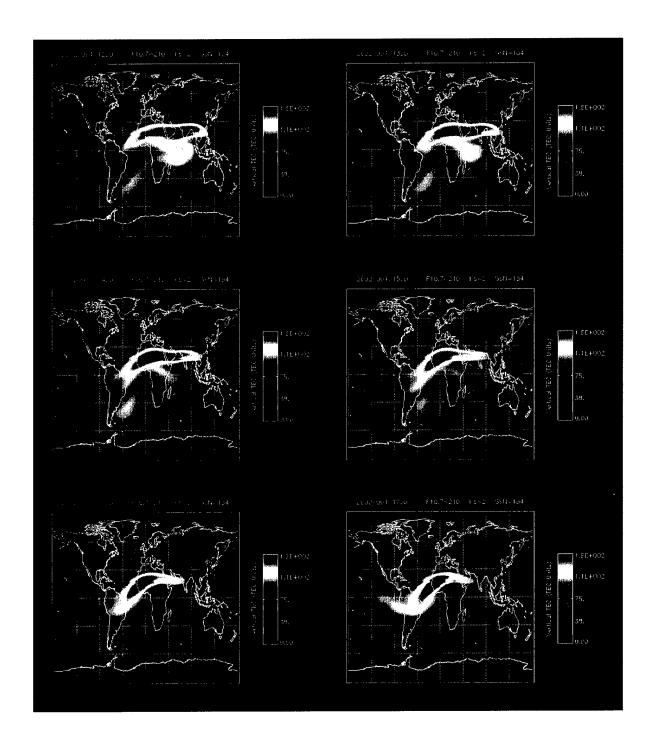
## Appendix F: Sample PRISM Output

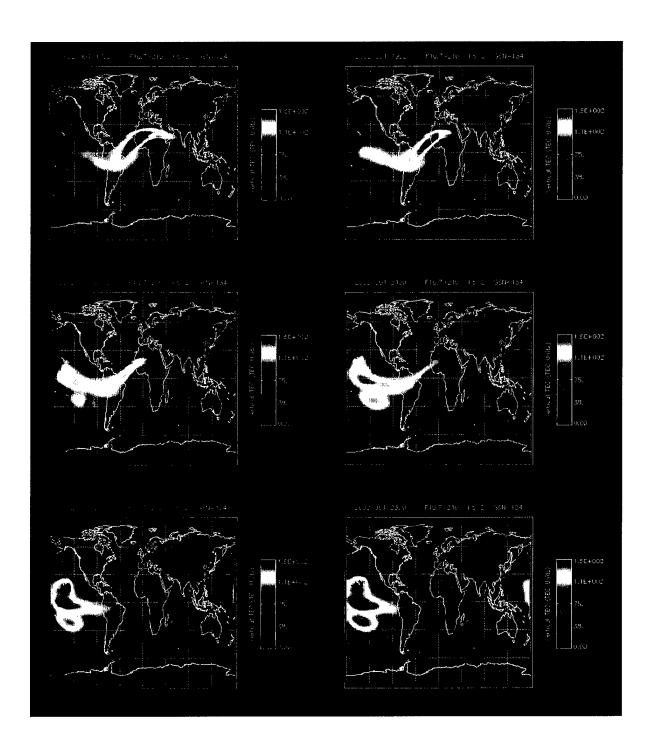
```
YEAR DAY UT (sec) F10.7 Kp Solar Sunspot Number
                     210.0 2.0 196.00
2002
       35
            0.0
 0
                           Latitude Longitude Latitude Longitude
   Latitude
              Longitude
 Starting Ending Starting Ending
                                  Step Step
                                                Delta Delta
                                                4.00
 -90.00 90.00 0.00 356.00
                              46
                                   90
                                           4.00
Number of altitude points =
 Altitudes
  90.00 95.00 100.00 105.00 110.00
 115.00 120.00 125.00 130.00 135.00
 140.00 145.00 150.00 160.00 170.00
 180.00 190.00 200.00 210.00 220.00
 230.00 240.00 250.00 260.00 270.00
 280.00 290.00 300.00 320.00 340.00
 360.00 380.00 400.00 450.00 500.00
 550.00 600.00 650.00 700.00 750.00
 800.00 850.00 900.00 1000.00 1100.00
 1200.00 1300.00 1400.00 1500.00 1600.00
-90.00 0.00 -74.50 17.30 20.28 CAP
Densities
7.83E+02 2.18E+03 6.05E+03 1.55E+04 2.50E+04
3.43E+04 4.43E+04 5.67E+04 7.09E+04 8.38E+04
9.54E+04 1.06E+05 1.16E+05 1.33E+05 1.48E+05
1.61E+05 1.73E+05 1.85E+05 1.97E+05 2.11E+05
2.26E+05 2.42E+05 2.60E+05 2.83E+05 3.07E+05
3.31E+05 3.53E+05 3.72E+05 3.92E+05 3.96E+05
3.84E+05 3.63E+05 3.37E+05 2.70E+05 2.12E+05
1.68E+05 1.33E+05 1.07E+05 8.77E+04 7.48E+04
6.37E+04 5.44E+04 4.64E+04 3.37E+04 2.45E+04
1.78E+04 1.30E+04 9.43E+03 6.86E+03 4.99E+03
 FoF2, HmF2, FoF1, HmF1, FoE, HmE TEC
 5.65 339.54 0.00 0.00 1.42 110.00 15.62
```

Appendix G: A sample of a PRISM TEC RTA output.









## References

- Anderson, D. N., A Theoretical study of the ionospheric F-Region equatorial anomaly, II, Results in the American and Asian sectors, Planet. *Space. Sci.*, 21, 421-442, 1973.
- Brace, L. H., and R. F. Theis, Global empirical models of ionospheric electron temperature in the upper F-region and plasmasphere based on in-situ measurements form the Atmosphere Explorer-C, ISIS 1, and ISIS 2 satellites, *J. Atmos. Terr. Phys.*, 43, 1317,1981.
- Callahan, P., TOPEX GDR user's handbook, JPL D-8944, Rev. A, internal document, pp.3-9, Jet Propulsion Lab, Pasadena, CA, 1993.
- Canck, Marcel H. De, Ionosphere The Earth's Atmosphere, AntenneX Online Issue No. 63, July 2002.
- Daniell, Robert E., L. D. Brown, Parameterized Real-Time Ionospheric Specification Model PRISM Version 1.5, Contract F19628-89-C-0005, Newton, MA, Computational Physics Inc., 31 Mar 1995.
- Decker, D. T., C. E. Valladares, R. Sheehan, Su. Basu, D. N. Anderson, and R. A. Heelis, Modeling daytime F layer patches over Sonderstrom, *Radio Science*, 29, 249-268, 1994.
- Hardy, D. A., M. S. Gussenhoven, R. R. Raistrick, and W. J. McNeil, Statistical and functional representation of the pattern of auroral energy flux, number flux, and conductivity, J. *Geophys. Res.*, 92 12275-12294, 1987.
- Hargreaves, J. K., The Solar-Terrestrial Environment, Cambridge, Cambridge University Press, 1992.
- Hedin, A. E., Empirical global model of upper thermosphere winds based on Atmospheric and Dynamics Explorer satellite data, *J. Geophys. Res.*, 93, 9959-9978, 1988.
- Heppner, J. P., and N. C. Maynard, Empirical high latitude electric field models, J. Geophys. Res., 92, 4467-4489, 1987.
- Ho, C. M., B.Wilson, A. Mannucci, U. Lindquister and D. Yuan, A Comparative Study of Ionospheric Total Electron Content Measurements Using Global

- Ionospheric Maps of GPS, TOPEX Radar, and the Bent Model, *Radio Science*, Vol 32, Number 4, pg 1499-1512, 1997.
- Imel, D. A., Evaluation of TOPEX/Poseidon dual-frequency ionosphere correction, J. Geophysics res., 99, 24895, 1994.
- Jasperse, J. R., The photoelectron distribution function in the terrestrial ionospherem in Physics of Space Plasmas, Scientific Publishers, Cambridge, MA, pp. 53-84, 1982.
- Jursa, Adolph S., Department of the Air Force, Handbook of Geophysics and the Space Environment, Air Force Geophysics lab, Air Force Systems Command, 1985.
- Mannucci, A., B. Wilson, and D. Yuan, Monitoring Ionospheric Total Electron Content Using the GPS Global Network and TOPEX/Poseidon Altimeter Data, Proceddings of the Beacon Satellite Symposium, University of Wales, Aberystwyth, 1994.
- Pulliam, R., W. Borer, D. Decker, P. Doherty, Operational Ionosphere Model Validation, Proceedings of the American Institute of Aeronautics and Astronautics Space 2000 Conference & Exposition, Paper #A00-42948, Long Beach, CA, Sep 2000.
- Rasinkangas, R. Textbook on Space Physics, International Space Physics Educational Consortium, November 1998.
- Strickland, D. J., D. L. Brook, T. P. Coffey, and J. A. Fedder, Transport equation techniques for the deposition of auroral electrons, *J. Geophysics Res.*, 81, 2755-2764, 1994.
- Tascione, Thomas F., Introduction to the Space Environment, 2<sup>nd</sup>Edition, Malabar, FL, Krieger Publishing, 1992.
- Vladimer, J. A., P. Jastrzebski, M. C. Lee, P. H. Doherty, D. T. Decker, D. N. Anderson, Longitude structure of ionospheric total electron Content at low latitudes measured by the TOPEX/Poseidon satellite, *Radio Science*, vol. 34, Number 5, 1239-1260, 1999.